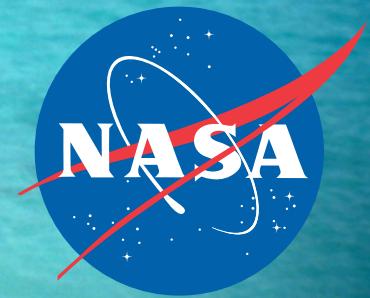




# *Trends in ocean carbon uptake*

Galen A. McKinley and Amanda Fay  
University of Wisconsin – Madison  
Atmospheric & Oceanic Sciences

NASA OCRT 2012  
Seattle, WA  
April 25, 2012



# Are trends in the global carbon cycle already detectable?

Reduction in fraction  
stored in the ocean  
(Canadell et al. 2007)

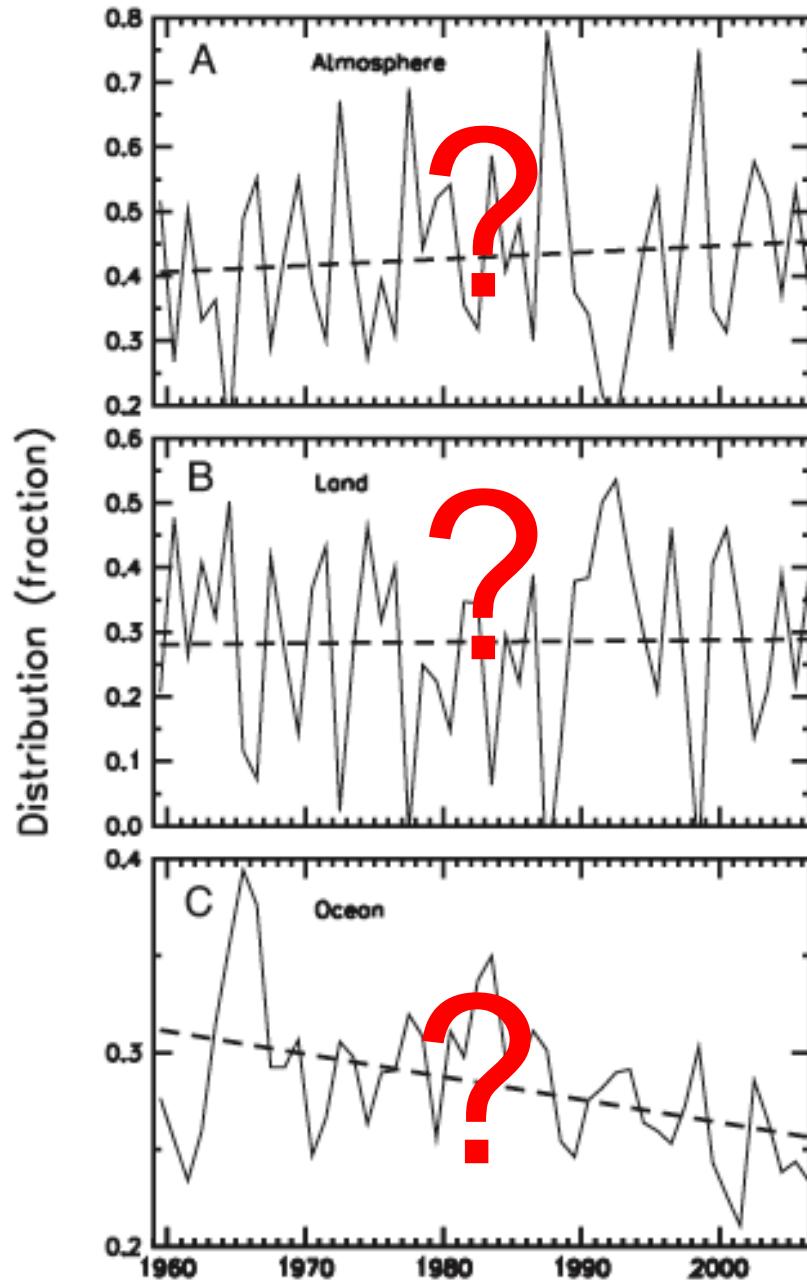


Fig. 2. Fraction of the total emissions ( $F_{\text{Foss}} + F_{\text{Uc}}$ ) that remains in the atmosphere (A), the land biosphere (B), and the ocean (C).

# Modeled-based assessment of global CO<sub>2</sub> sink impacts due to climate feedbacks: 1981-2007 = -0.20 PgC/decade

Mechanism	Sink impact	Regional notes
Warming	-20%	50% in North Atlantic alone
Wind change	-63%	>80% in Tropical Pacific >30% in S. Ocean Compensation elsewhere
Heat, Freshwater flux	+15%	In Northern Hemisphere
Nonlinear	-32%	>65% in Tropics

LeQuéré et al. 2010

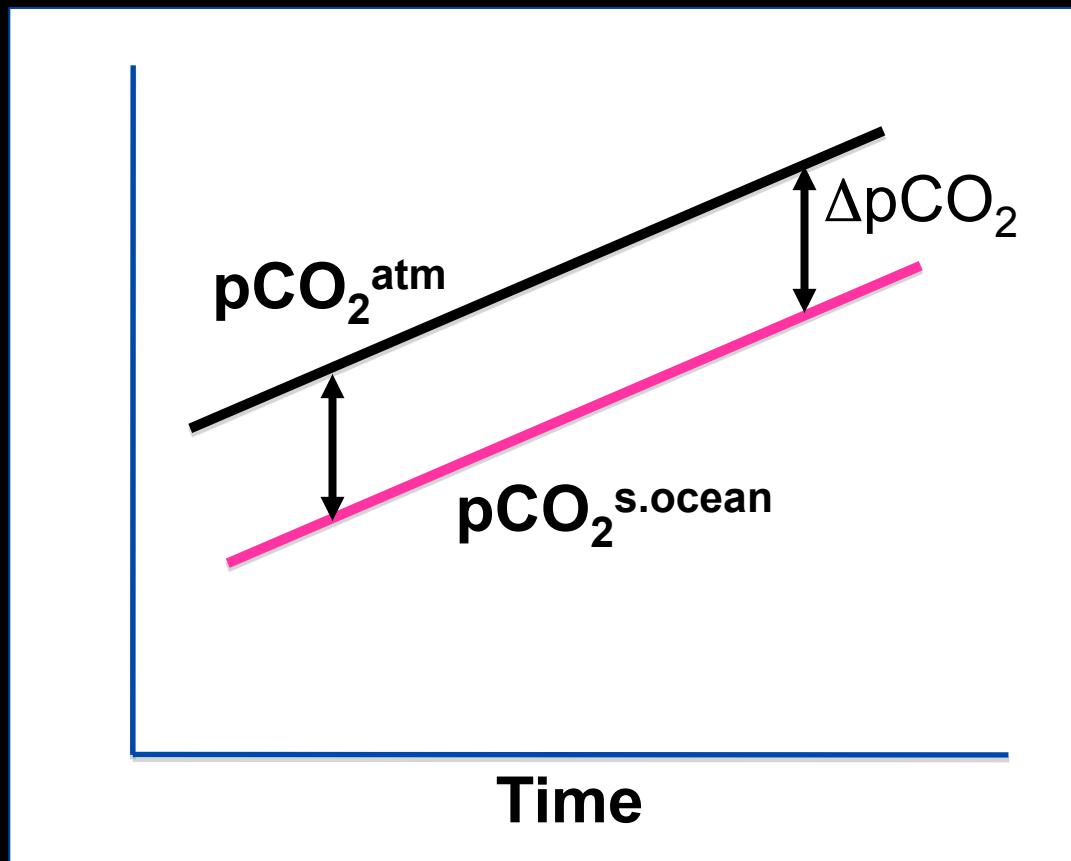
Earlier modeling found that warming to be dominant negative feedback (-84%) on CO<sub>2</sub> sink in first 100yrs after 4xCO<sub>2</sub> pulse (Sarmiento and LeQuéré, 1996)

# Outline

- Background: Ocean carbon sink trends from surface ocean pCO<sub>2</sub>
- N. Atlantic trends from Takahashi database
  - Gyre-scale biomes
  - Distinguishing variability from trends
  - Subtropics: Negative feedback from warming
- Global extension

# Trends from surface ocean $p\text{CO}_2$ :

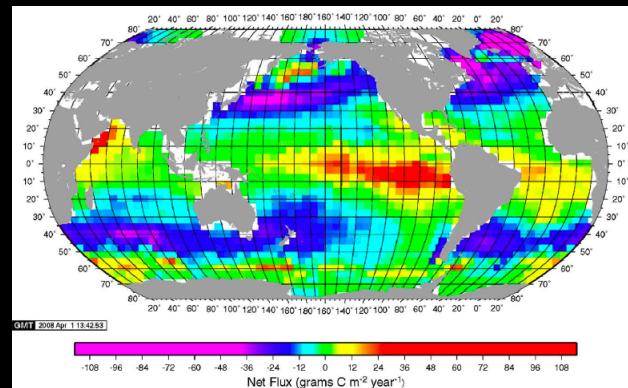
If  $p\text{CO}_2^{\text{atm}}$  only change, i.e. circulation, biology constant



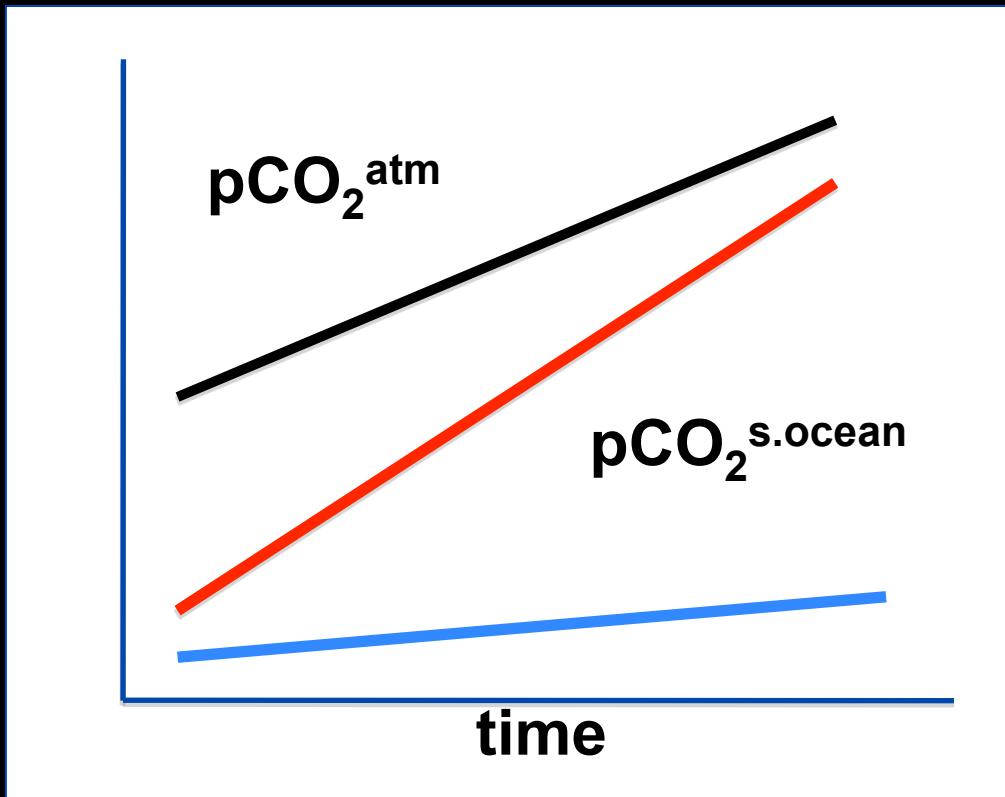
“Equilibration”  
 $d\text{pCO}_2^{\text{s.ocean}}/dt = d\text{pCO}_2^{\text{atm}}/dt$

$$d\Delta p\text{CO}_2/dt = 0$$
$$d(\text{CO}_2 \text{Flux})/dt = 0$$

STEADY SINKS AND SOURCES



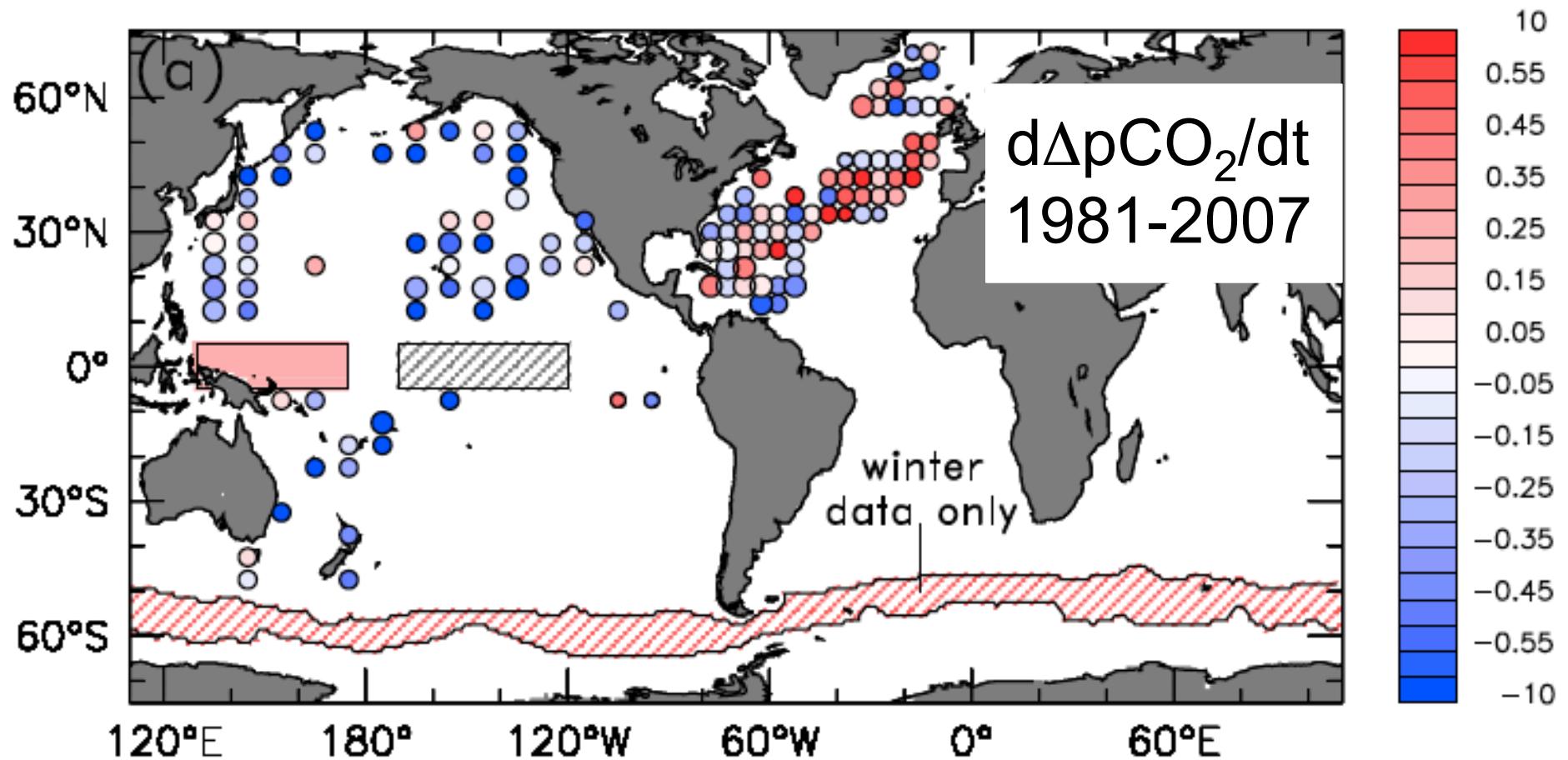
As corollary,  $\text{dpCO}_2^{\text{s.ocean}}/\text{dt} \neq \text{dpCO}_2^{\text{atm}}/\text{dt}$   
has been interpreted as a change in flux  
*due to change in biology or circulation*



$\text{dpCO}_2^{\text{s.ocean}}/\text{dt} >$   
 $\text{dpCO}_2^{\text{atm}}/\text{dt}$   
**DECLINING SINK**

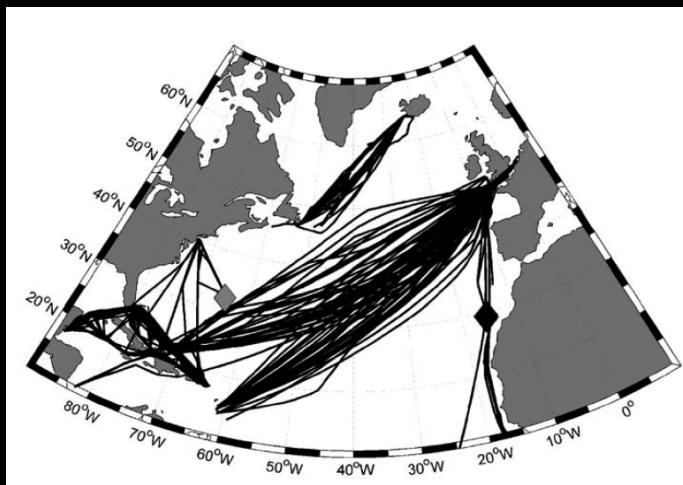
$\text{dpCO}_2^{\text{s.ocean}}/\text{dt} <$   
 $\text{dpCO}_2^{\text{atm}}/\text{dt}$   
**INCREASING SINK**

# Mostly from Takahashi pCO<sub>2</sub> database, model in S. Ocean

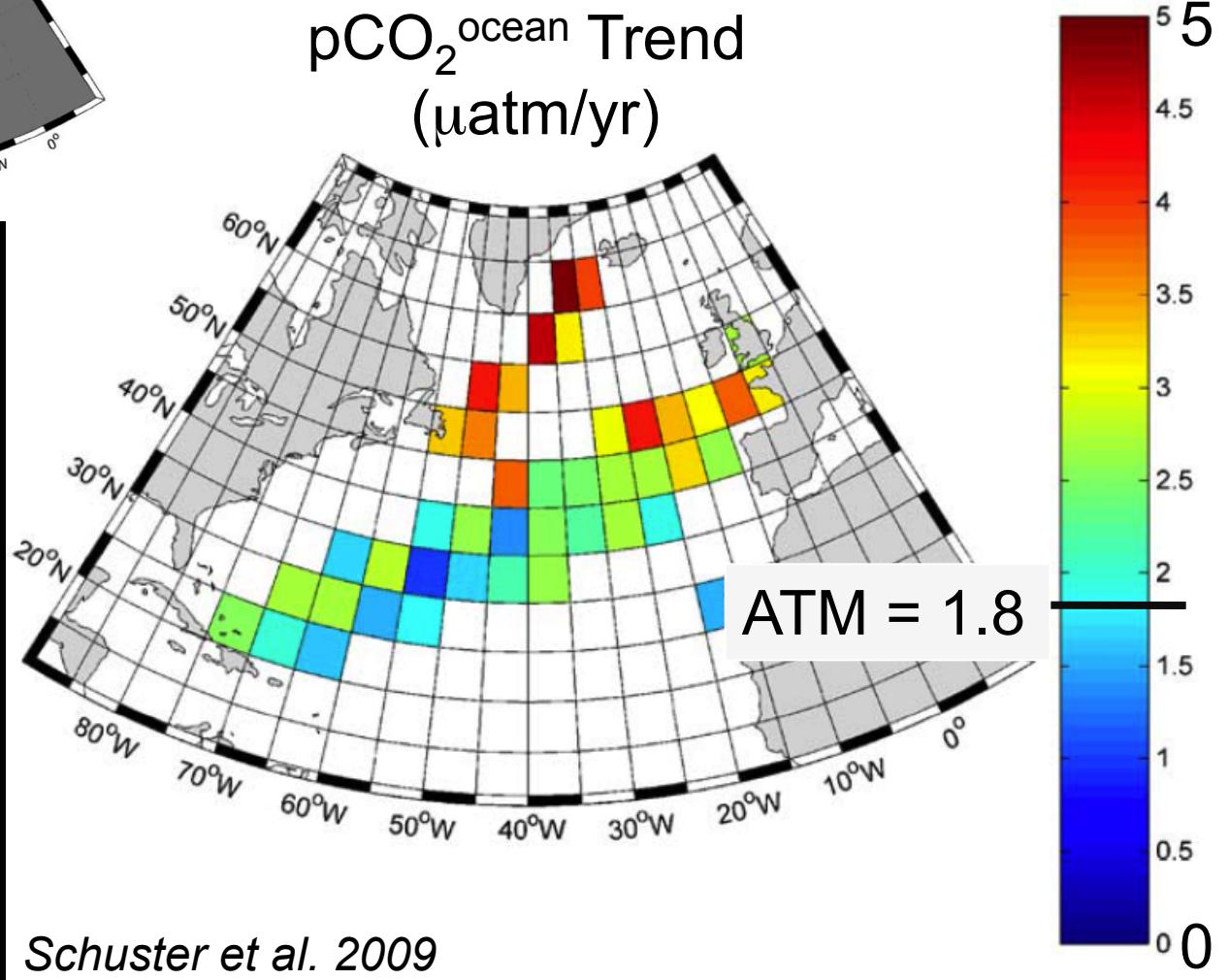


LeQuéré et al. 2010

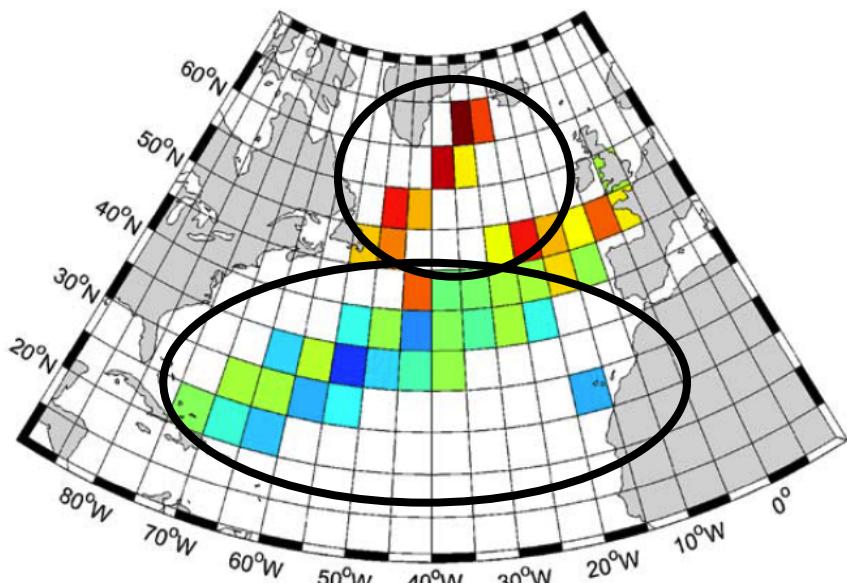
# N. Atlantic: VOS datasets, linear trend 1990-2006



Data of  
Corbiere et al. 2007  
Shuster & Watson 2007  
Bates 2007  
Olsen et al. 2004  
Santana-Casiano et al. 2007

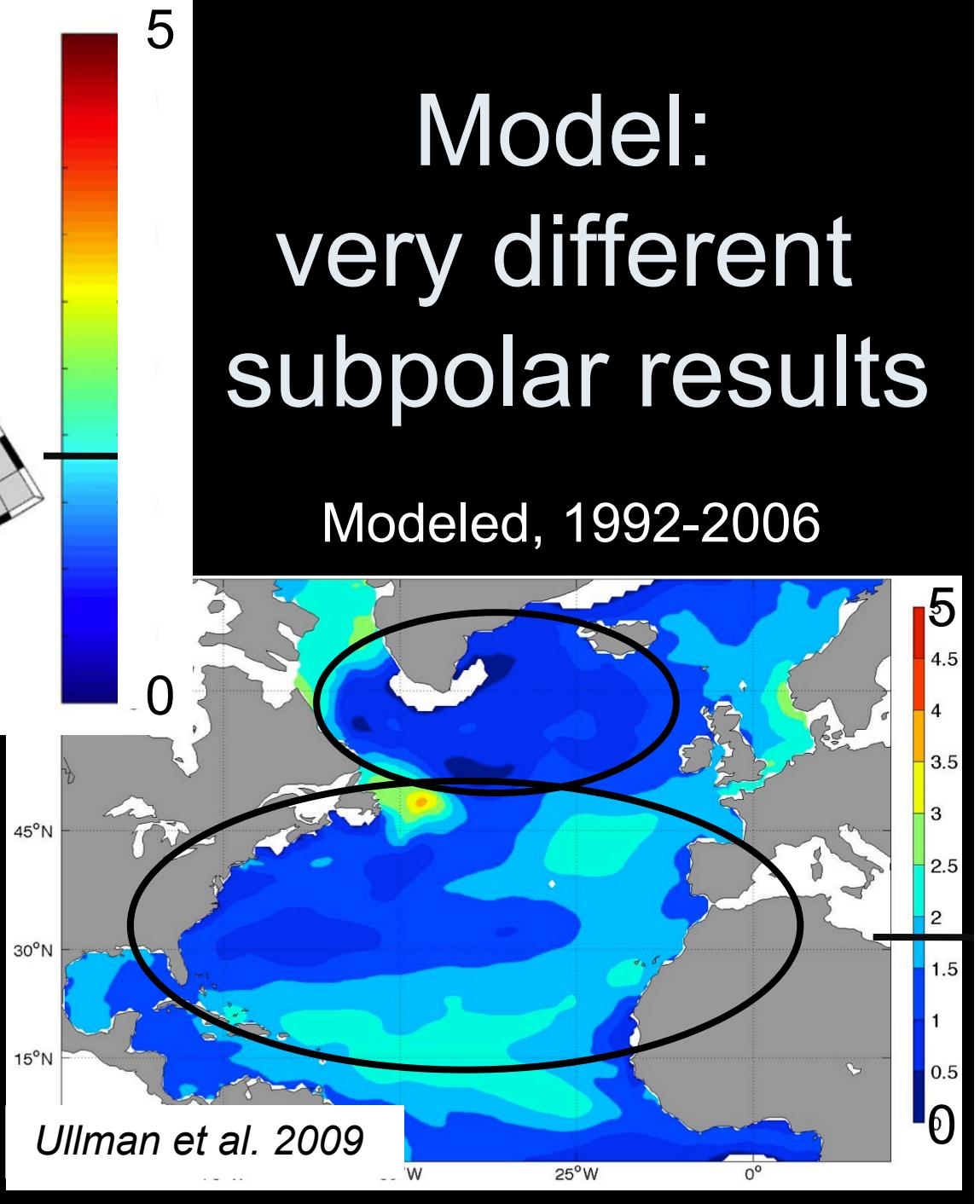


Observed 1990-2006



*Schuster et al. 2009*

Generally consistent  
<45N, but  
inconsistent >45N



# Observed N. Atlantic pCO<sub>2</sub> trends

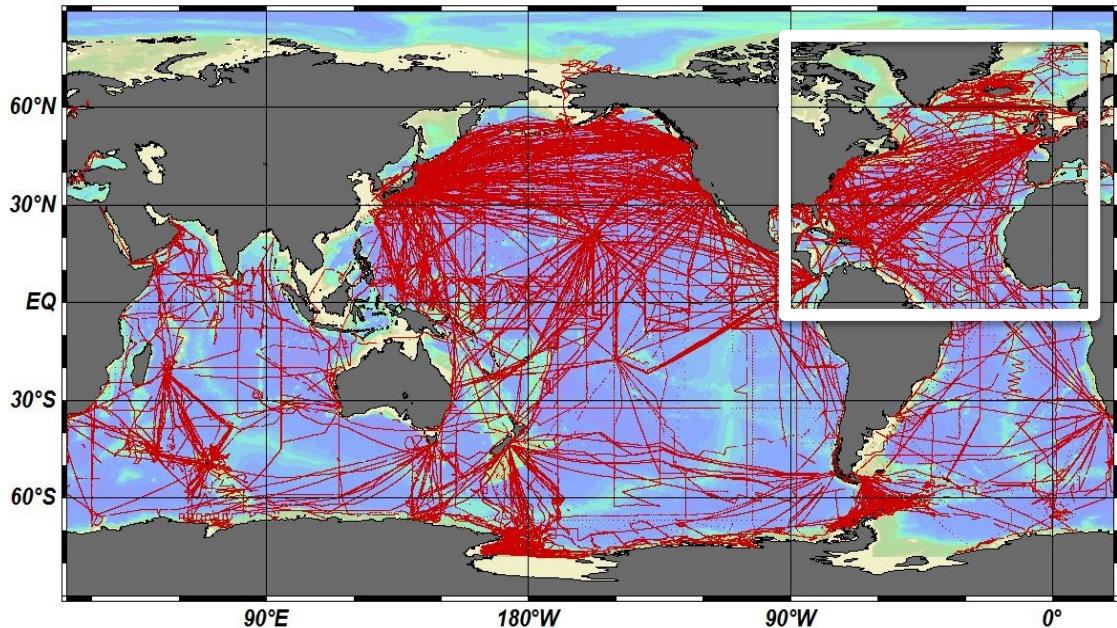
*McKinley et al. 2011*

# Objectives

with *in situ* pCO<sub>2</sub> data, discern:

- Timescales: variability vs. trends
- Mechanisms: Warming vs. carbon uptake
- Spatial scale: Gyre-size “biomes”

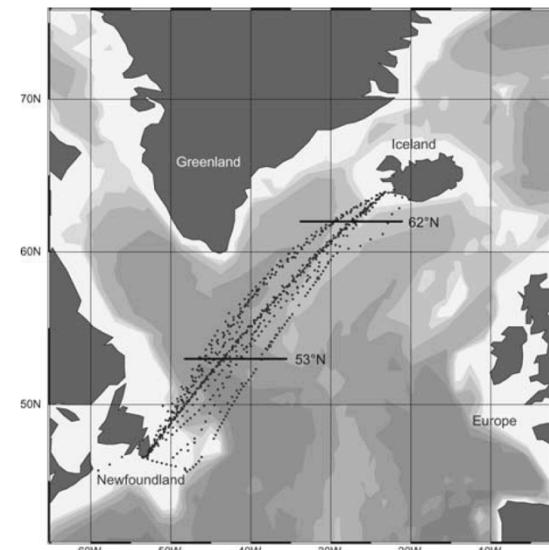
# Takahashi et al. 2010 in-situ pCO<sub>2</sub> database released on CDIAC website



SURATLANT dataset  
Calculated pCO<sub>2</sub>  
(1993-2007)

-Over 4.5 million data points globally

-Over 1 million in the North Atlantic



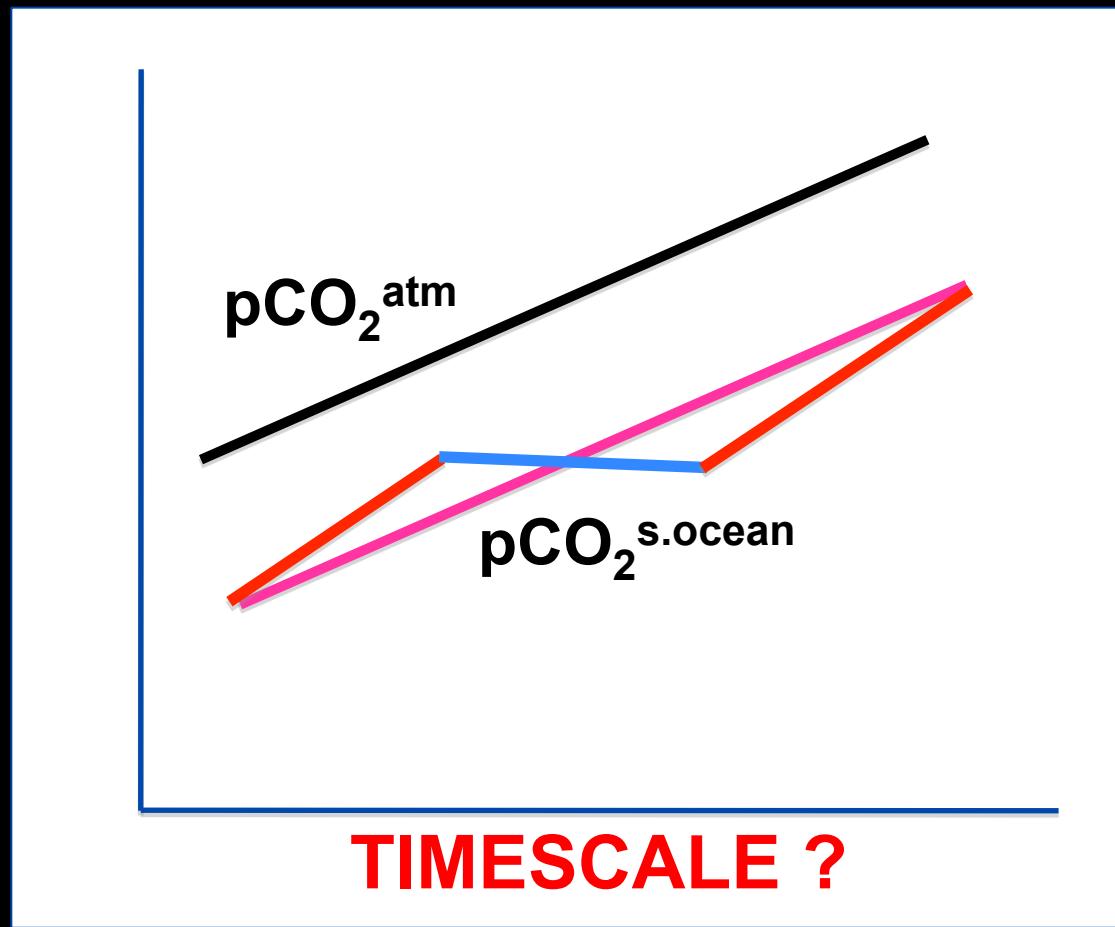
Corbiere et al. 2007

# Objectives

with *in situ* pCO<sub>2</sub> data, discern:

- Timescales: variability vs. trends
- Mechanisms: Warming vs. carbon uptake
- Spatial scale: Gyre-size “biomes”

# On what timescale does the ocean follow the atmosphere?

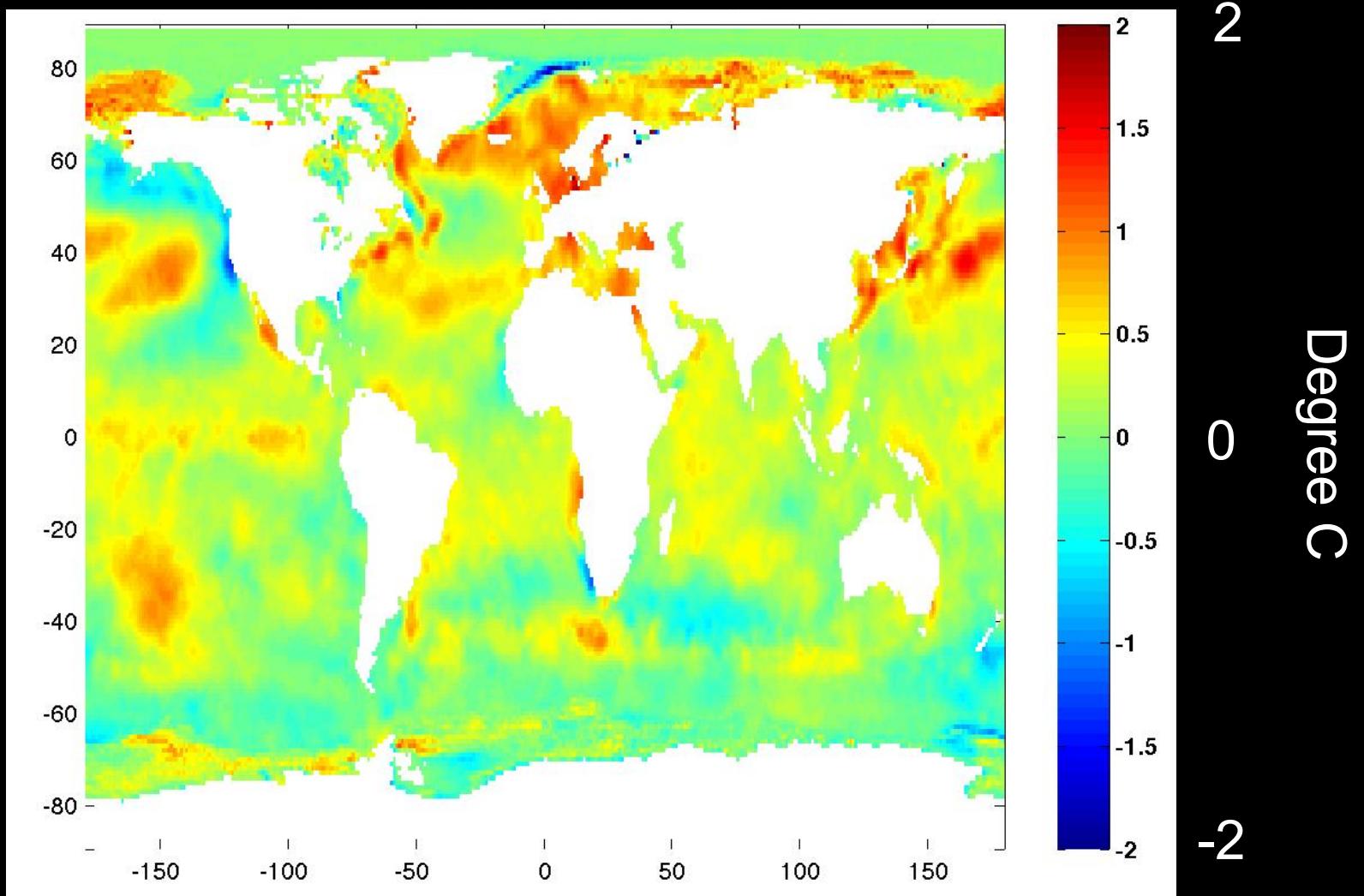


# Objectives

with *in situ* pCO<sub>2</sub> data, discern:

- Timescales: variability vs. trends
- Mechanisms: Warming vs. carbon uptake
- Spatial scale: Gyre-size “biomes”

# Observed SST change (2000-09) – (1980-89)



HadISST; Rayner et al 2003

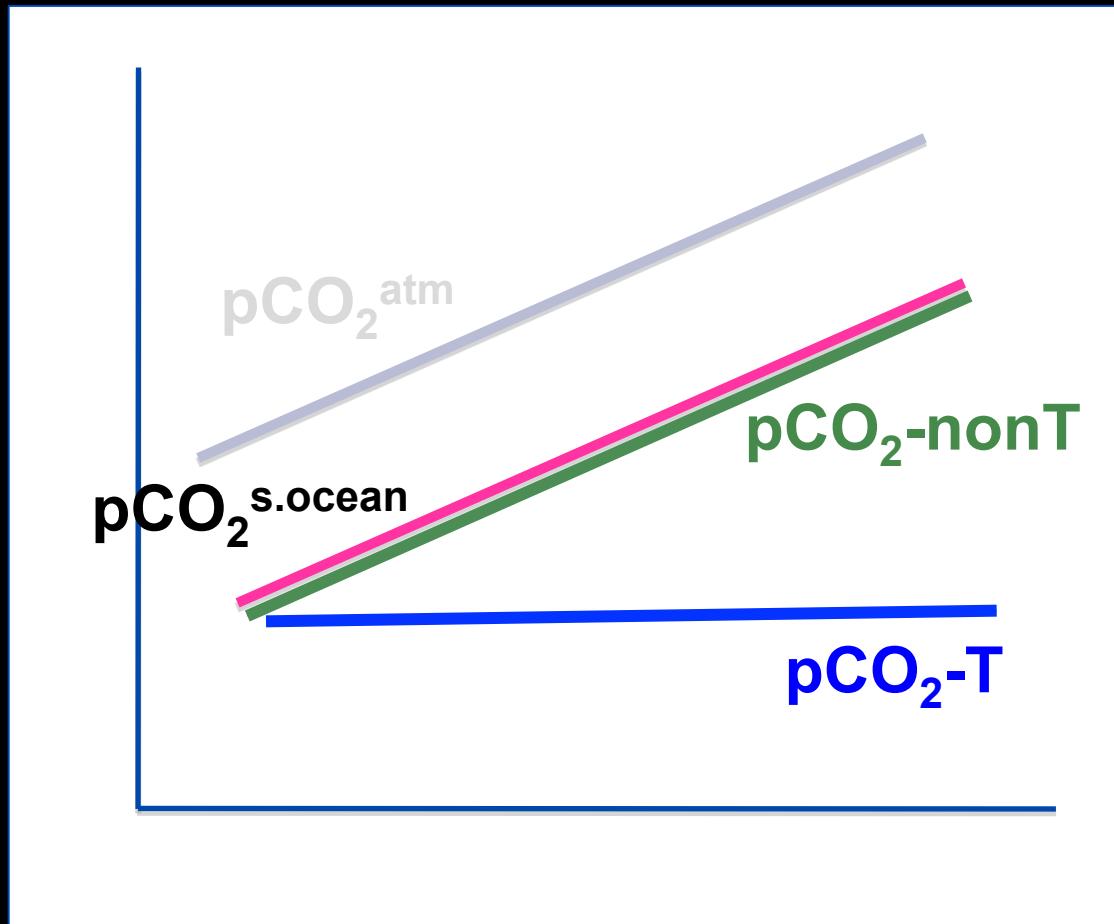
Decompose  $pCO_2$  into temperature driven component ( $pCO_2\text{-}T$ ) and biological/chemical component ( $pCO_2\text{-}nonT$ )

$$pCO_2 - T = \overline{pCO_2} * \exp(0.0423 * (SST - \overline{SST}))$$

$$pCO_2 - nonT = pCO_2 * \exp(0.0423 * (\overline{SST} - SST))$$

*Takahashi et al. 2002*

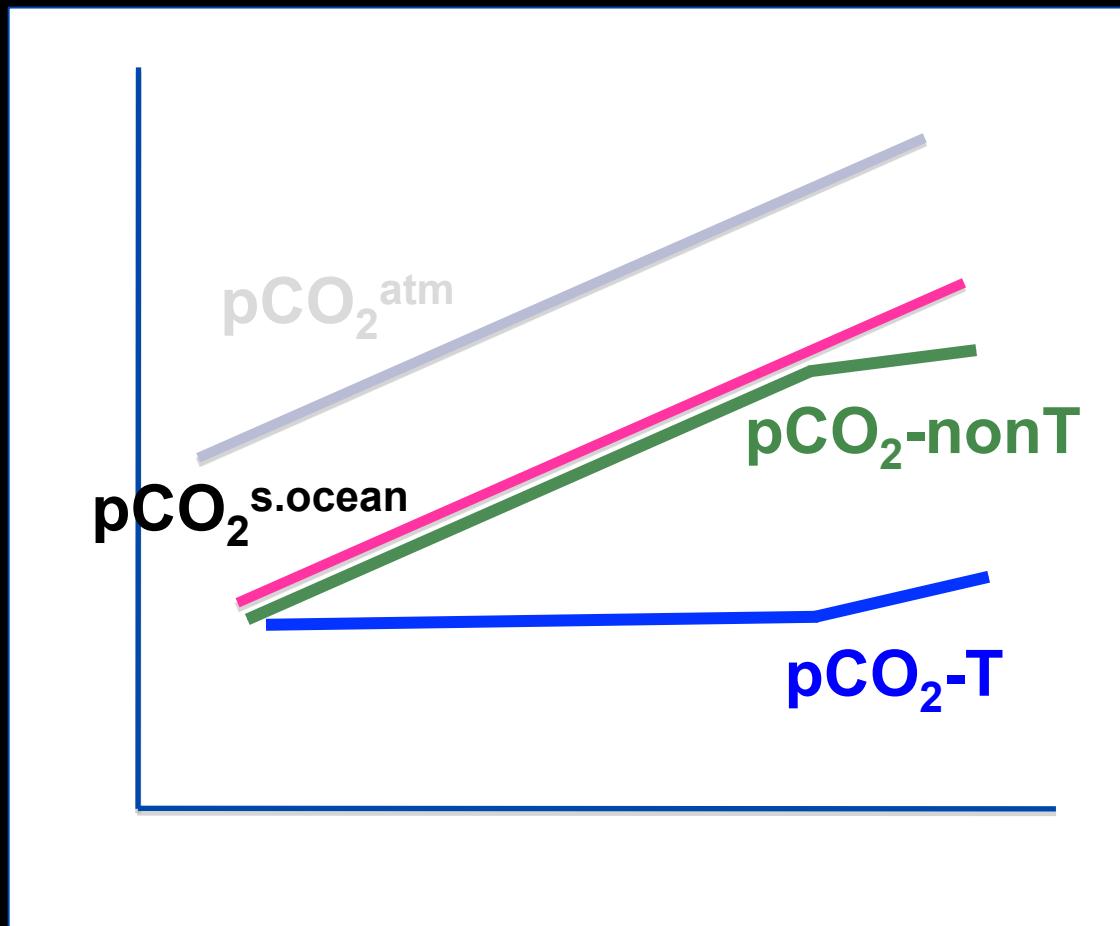
# Trend mechanisms



Biogeochemical  
change only

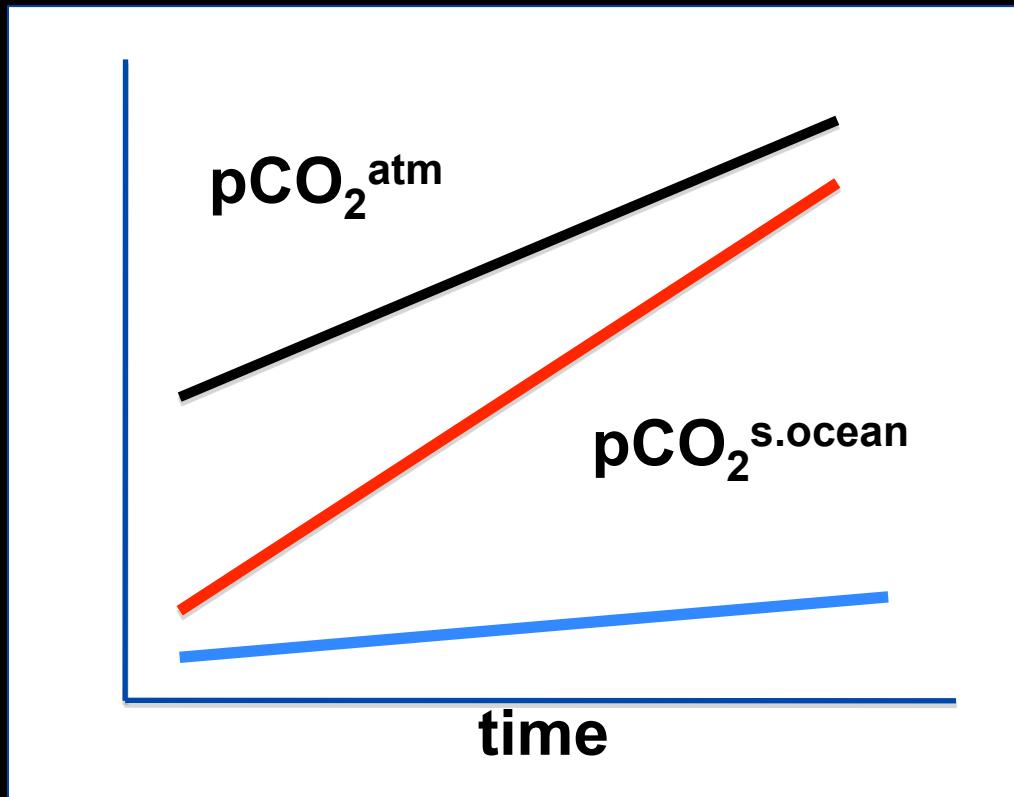
Consistent with  
carbon uptake

# Trend mechanisms



If warming contributes, less carbon uptake required for equilibration with atmosphere

# Mechanistic approach allows more nuanced understanding



$$\frac{dpCO_2^{s.ocean}}{dt} > \frac{dpCO_2^{atm}}{dt}$$

**OVER-EQUILIBRATION**

$$\frac{dpCO_2^{s.ocean}}{dt} < \frac{dpCO_2^{atm}}{dt}$$

**UNDER-EQUILIBRATION**

# Objectives

## with *in situ* pCO<sub>2</sub> data, discern:

- Timescales: variability vs. trends
- Mechanisms: Warming vs. carbon uptake
- Spatial scale: Gyre-size “biomes”

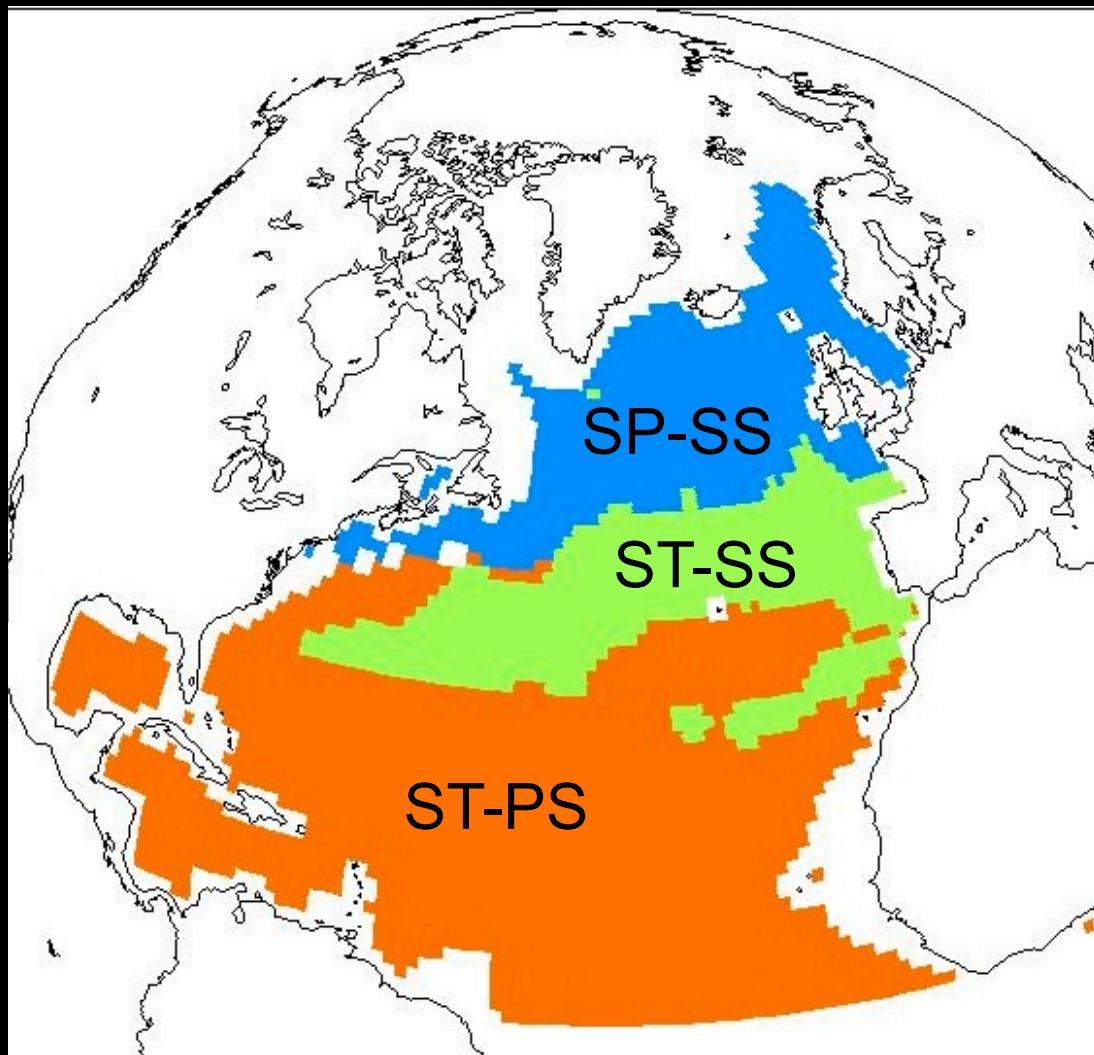
# Biomes

Seasonally  
Stratified  
Subpolar  
(SP-SS)

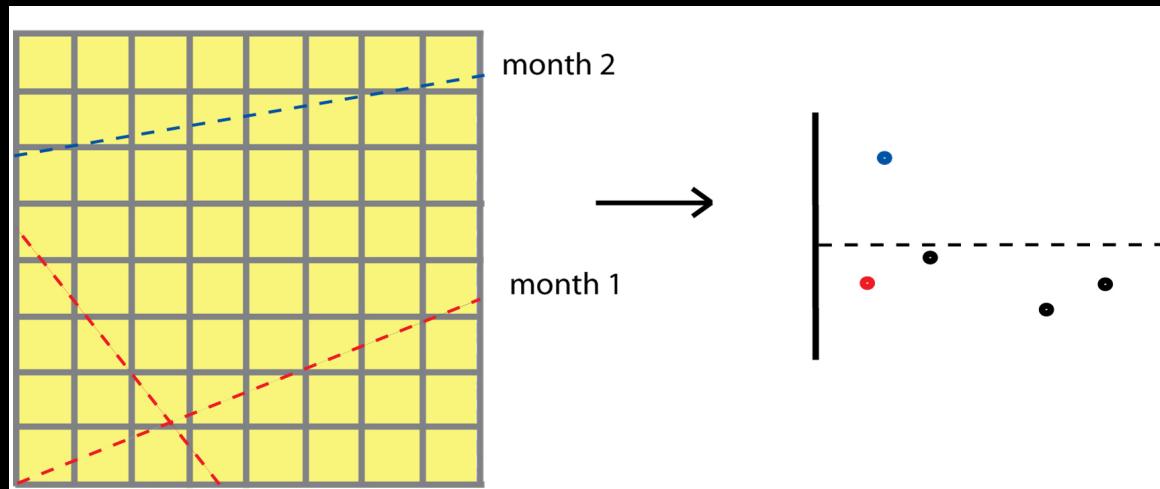
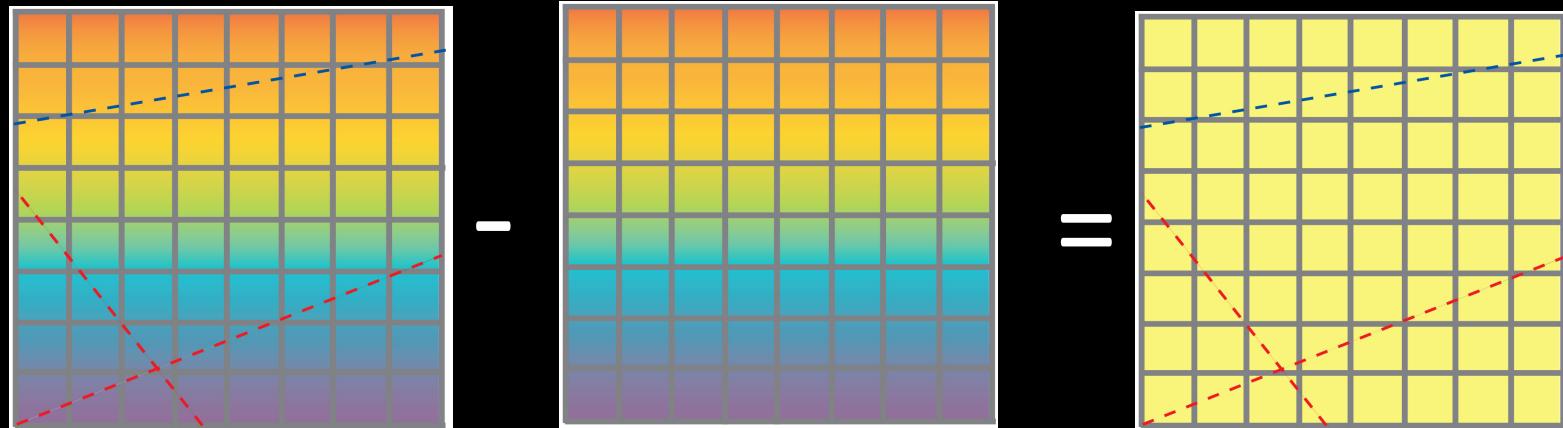
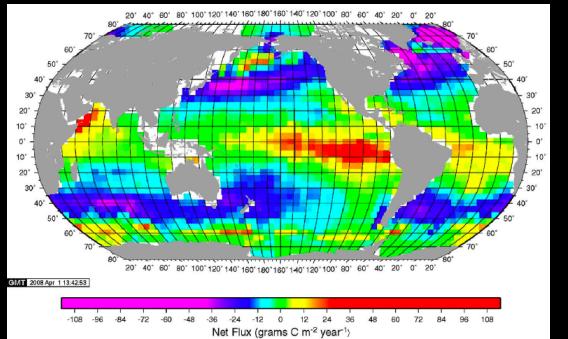
Seasonally  
Stratified  
Subtropical  
(ST-SS)

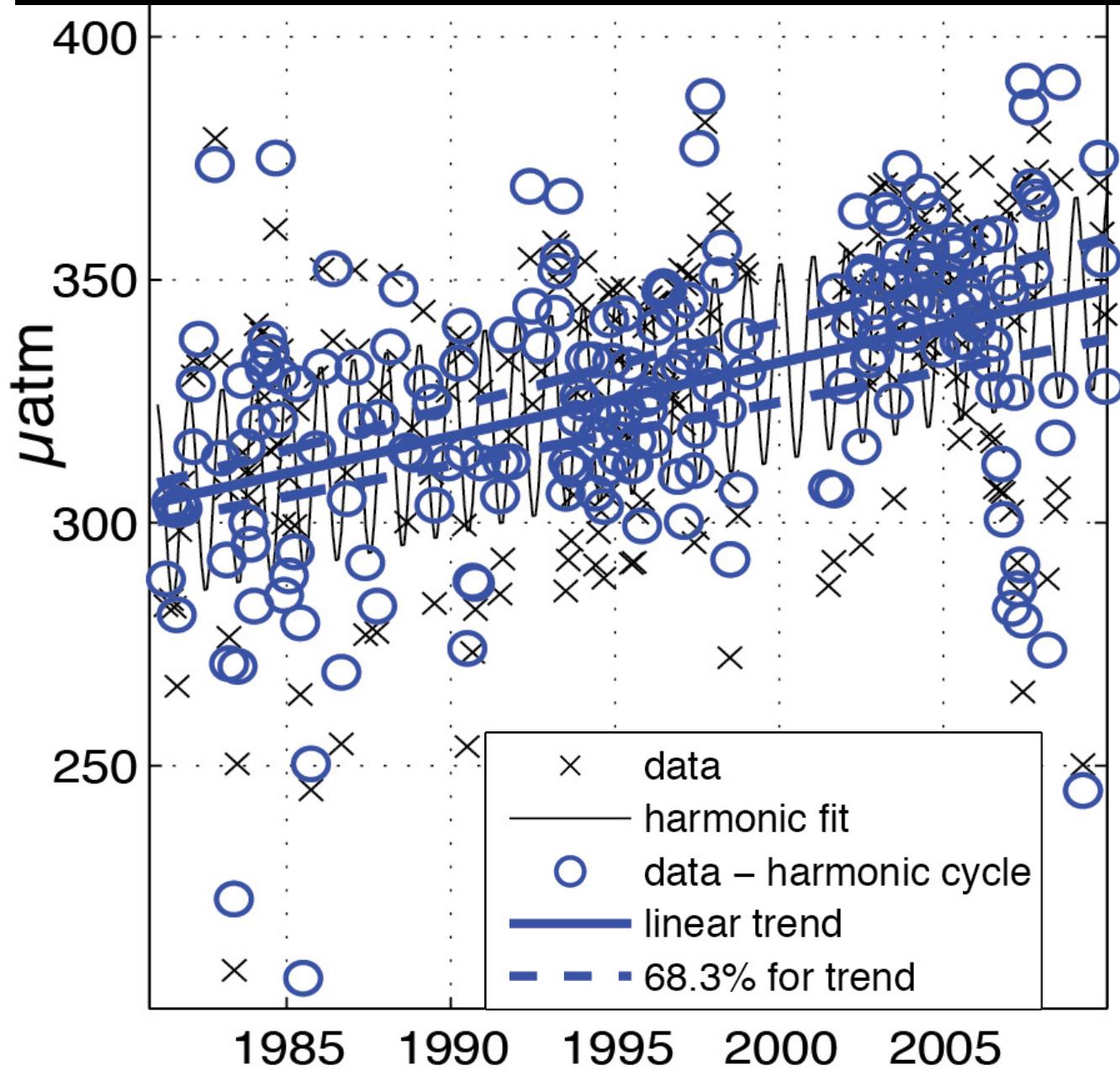
Permanently  
Stratified  
Subtropical  
(ST-PS)

- Gyre-scale biogeographic regions (Sarmiento et al (2004))
- Selection criteria: observed climatological SST, max MLD, and chlorophyll-a

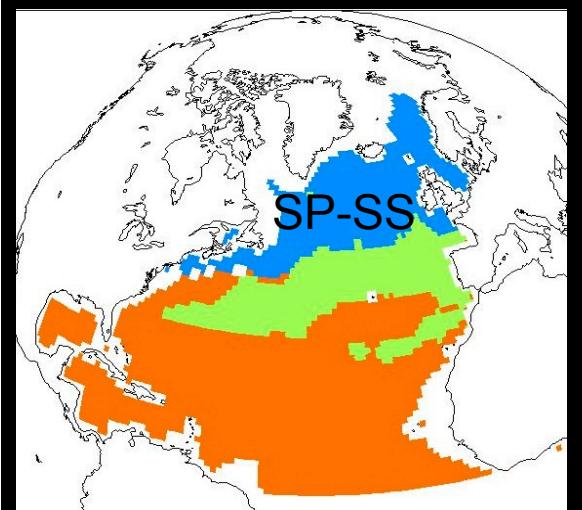


1. Subtract background mean to address spatial aliasing
2. Average to biome timeseries

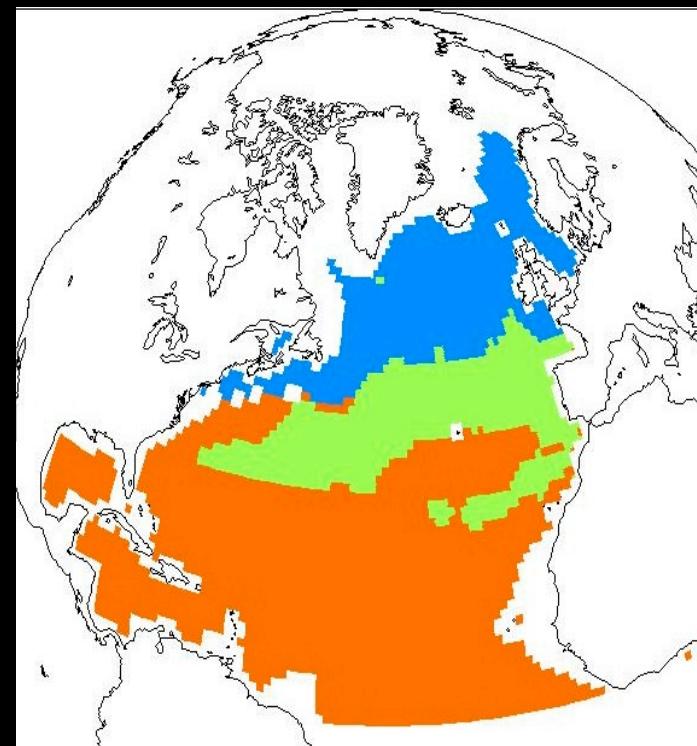
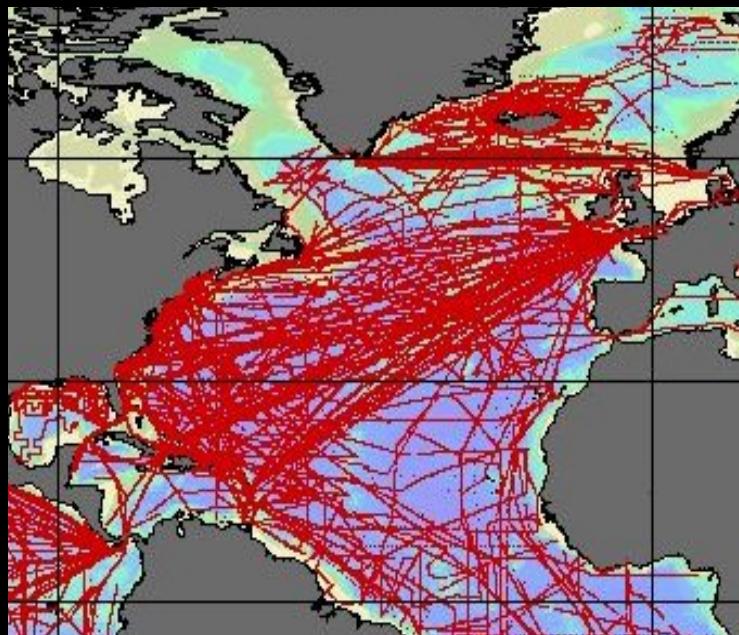




Resulting  $\text{pCO}_2$   
timeseries,  
subpolar  
N. Atlantic



# How representative are these trends of the biome?



Real World

Model

?

True

pCO<sub>2</sub>  
observations

Sampled

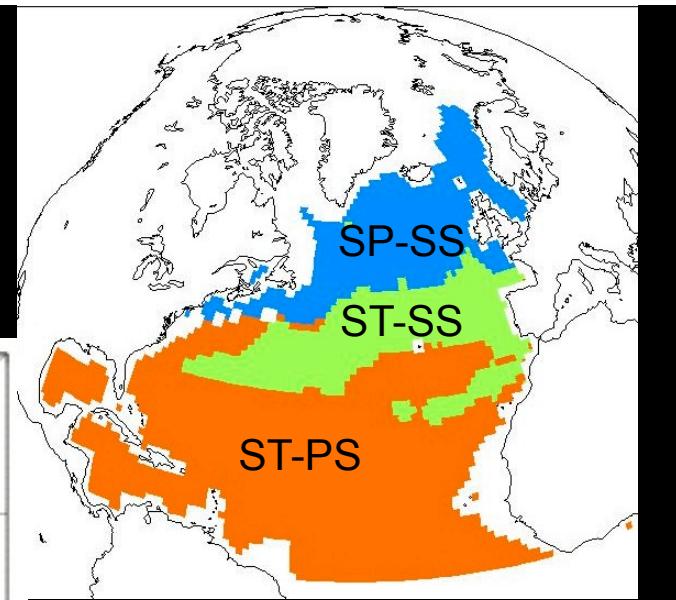
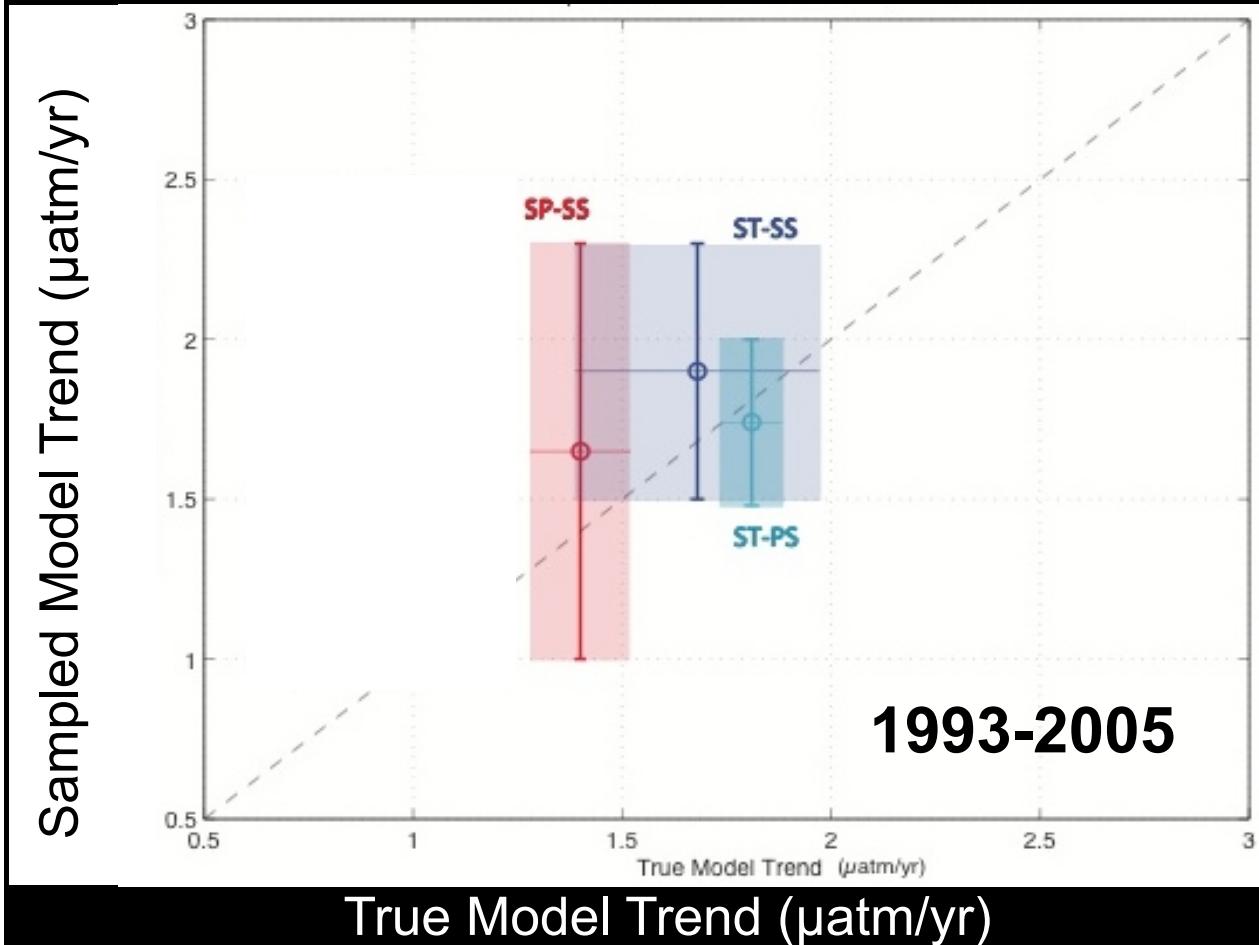
Biogeochemical  
model (MITgcm.NA)

True

Model  
subsampled  
as data

Sampled

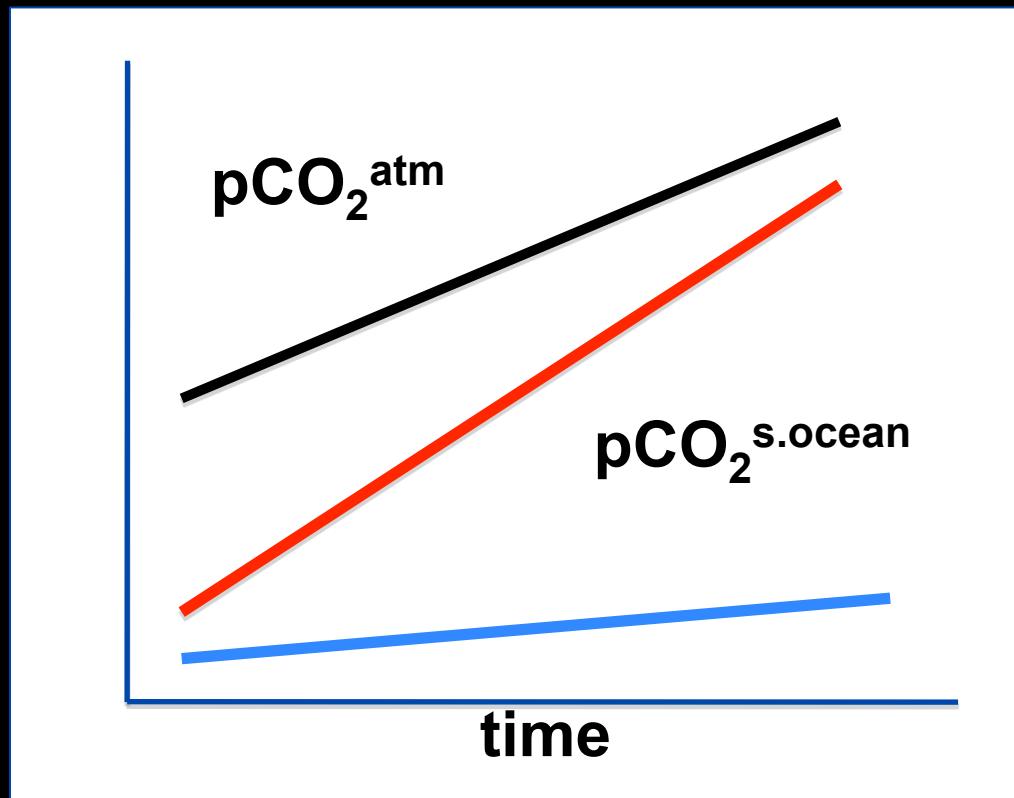
# Model Confirmation



Sampled, analyzed model captures trends from all model output at  $1\sigma$  level

Indicates that sampling and analysis technique are representative

# Results



$\frac{dp\text{CO}_2^{\text{s.ocean}}}{dt} > \frac{dp\text{CO}_2^{\text{atm}}}{dt}$

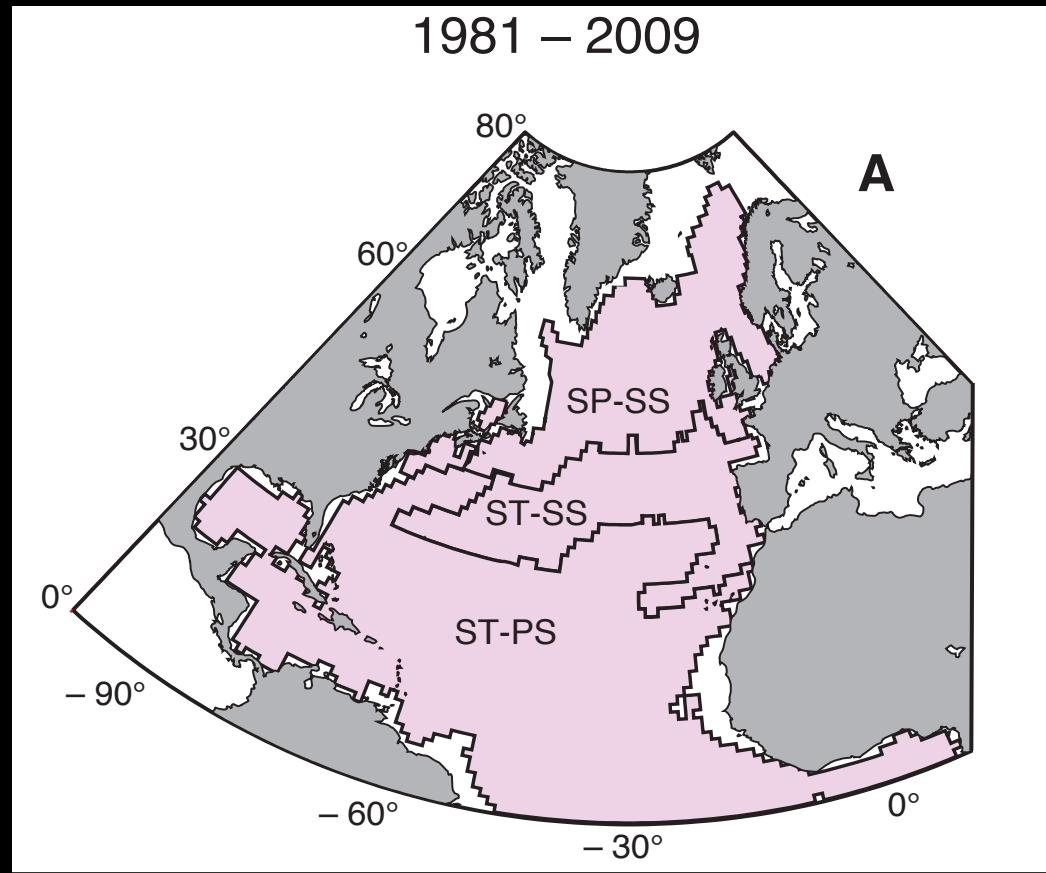
**OVER-EQUILIBRATION**

$\frac{dp\text{CO}_2^{\text{s.ocean}}}{dt} < \frac{dp\text{CO}_2^{\text{atm}}}{dt}$

**UNDER-EQUILIBRATION**

Multi-decadal:  
1981-2009

## Trend in $pCO_2^{ocean}$ compared to $pCO_2^{atm}$



$$\frac{dpCO_2^{ocn}}{dt} < \frac{dpCO_2^{atm}}{dt}$$

under-equilibration

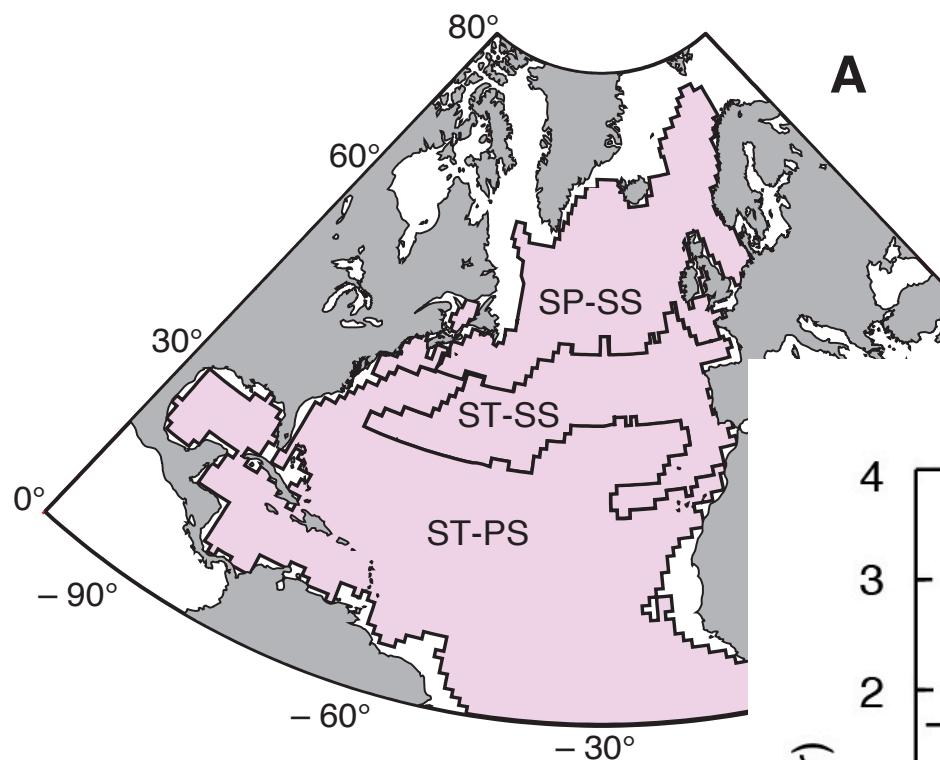
$$\frac{dpCO_2^{ocn}}{dt} \sim \frac{dpCO_2^{atm}}{dt}$$

equilibration

$$\frac{dpCO_2^{ocn}}{dt} > \frac{dpCO_2^{atm}}{dt}$$

over-equilibration

1981 – 2009

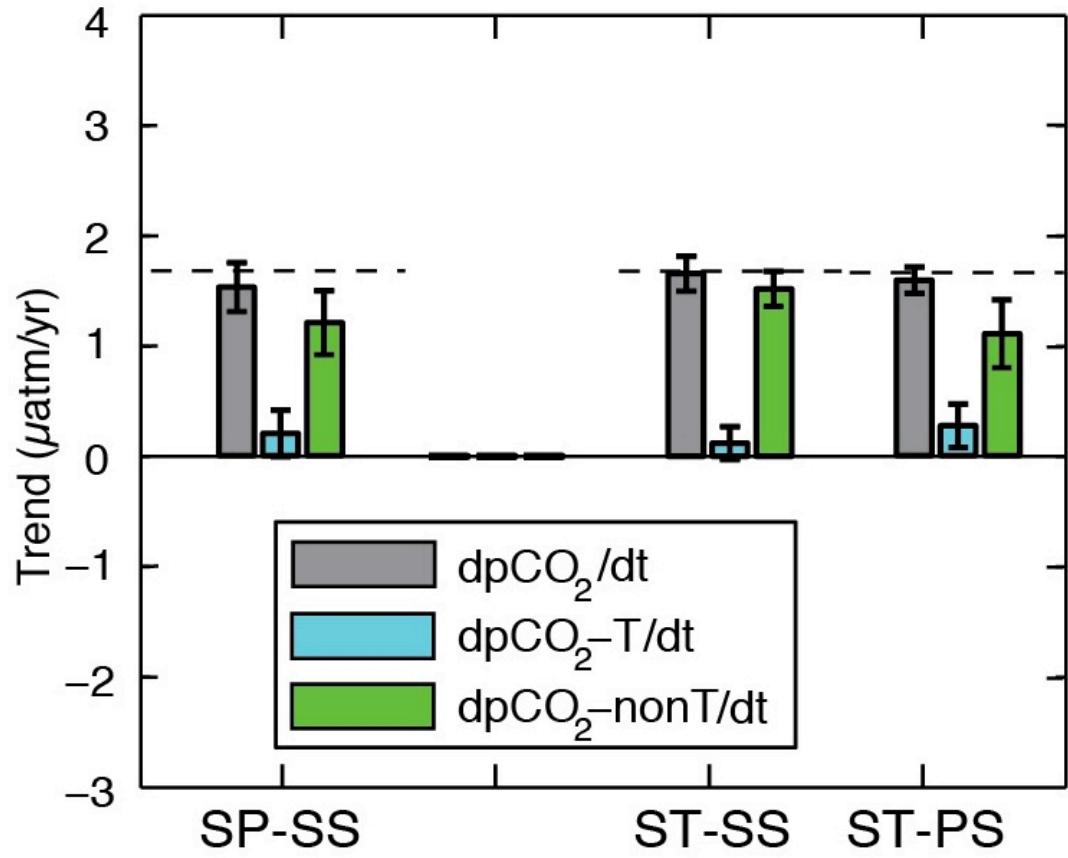


A

Chemical change  
dominates on  
multidecadal  
timescale

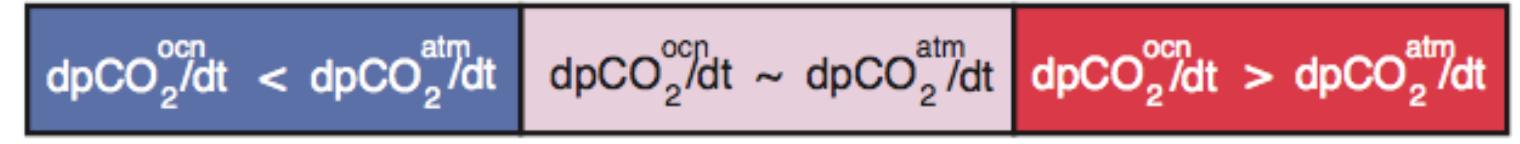
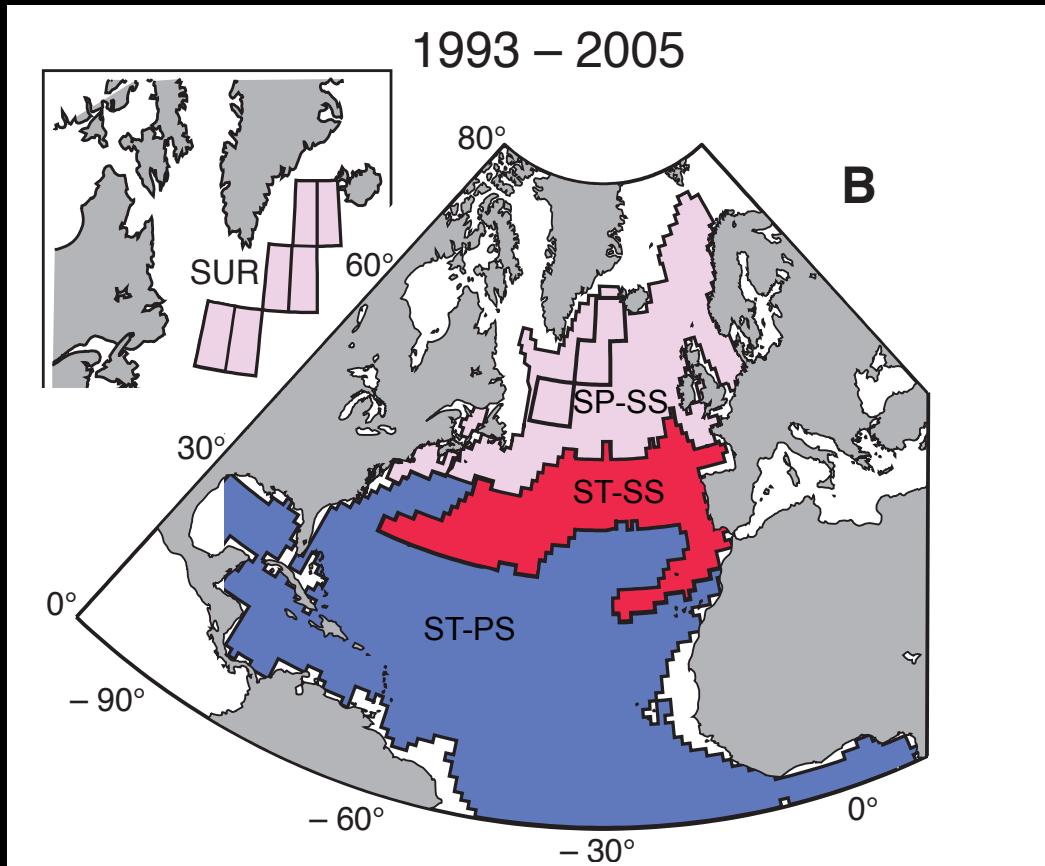
c

1981 – 2009



Decadal:  
1993-2005

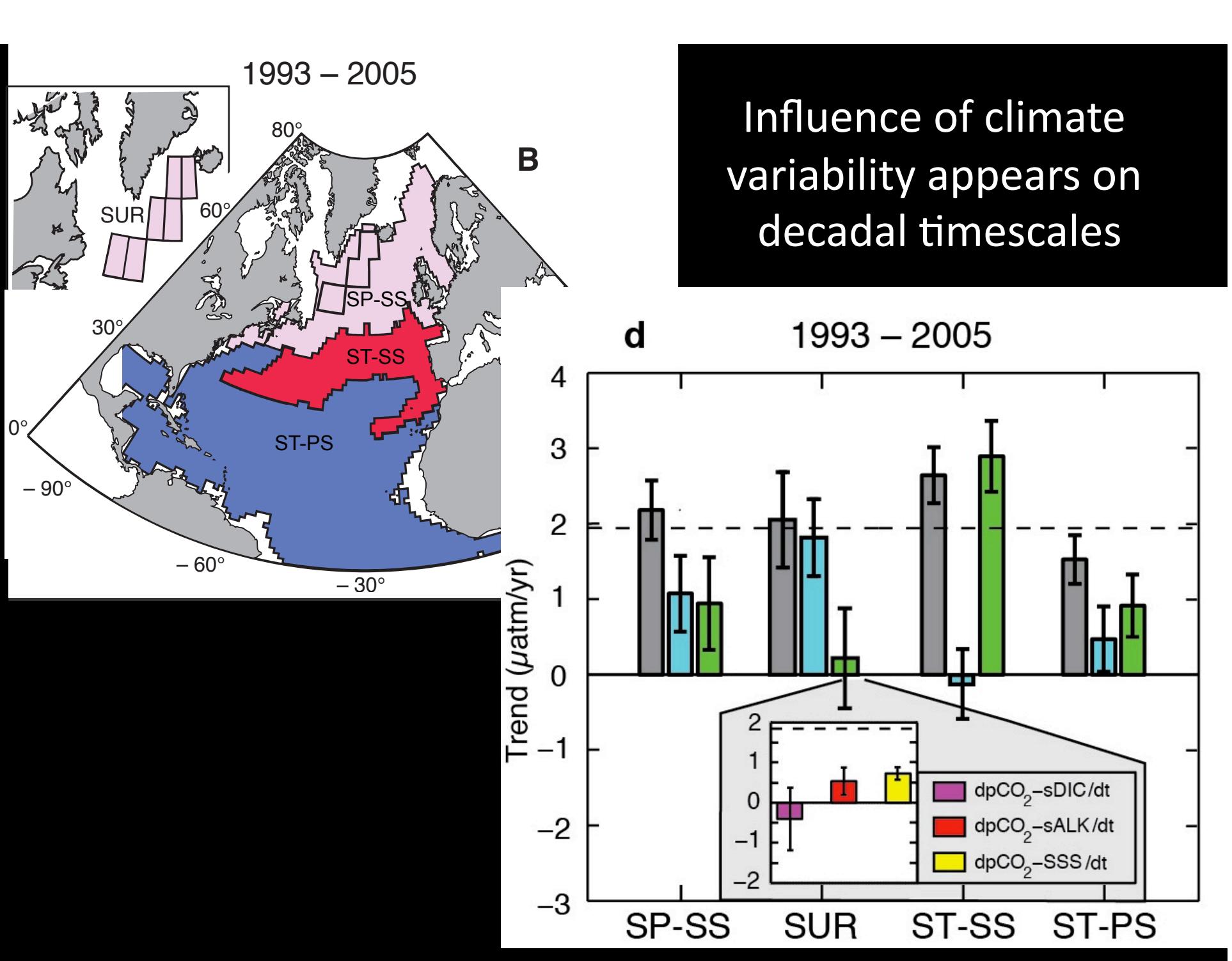
## Trend in $pCO_2^{ocean}$ compared to $pCO_2^{atm}$



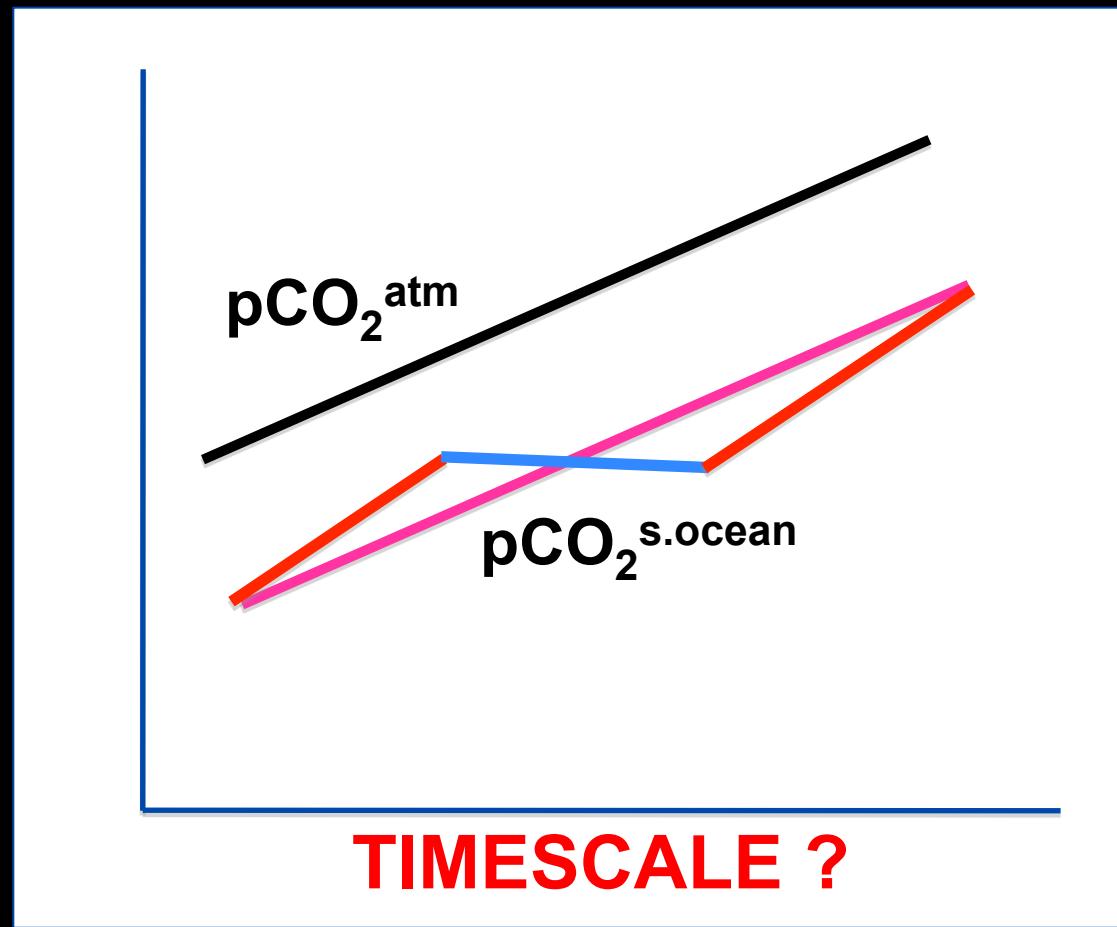
under-equilibration

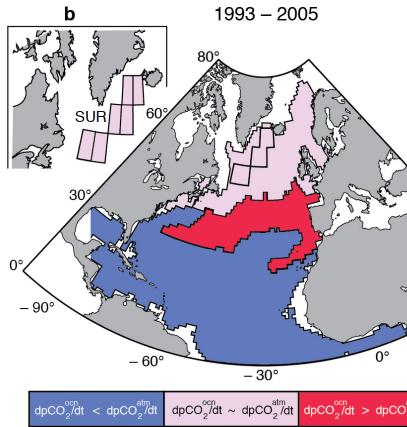
equilibration

over-equilibration

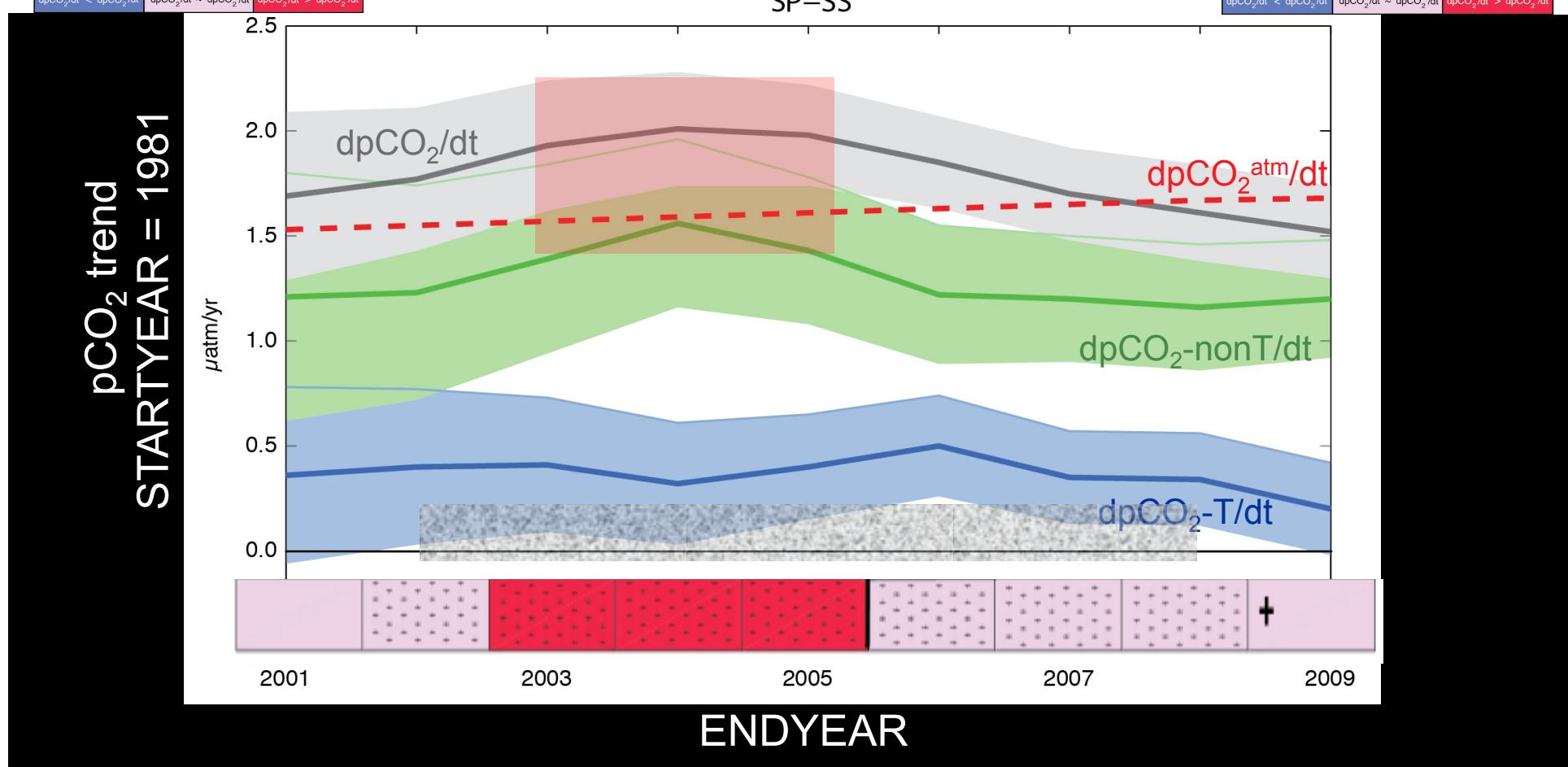
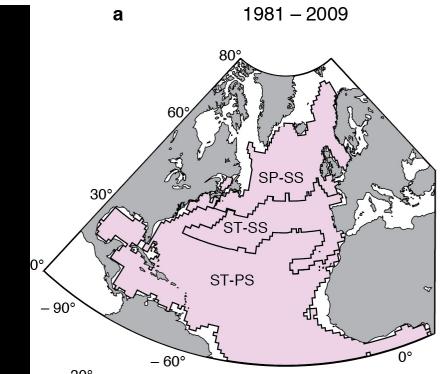


# On what timescale does the ocean follow the atmosphere?



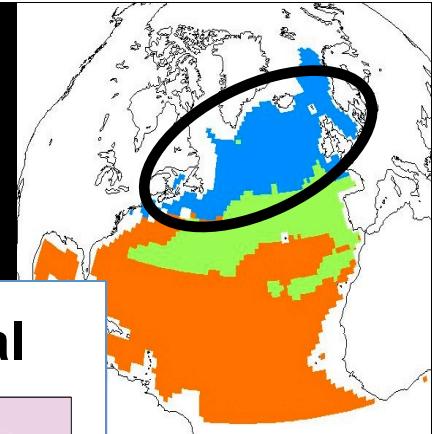


# Evaluate transition from decadal to multidecadal



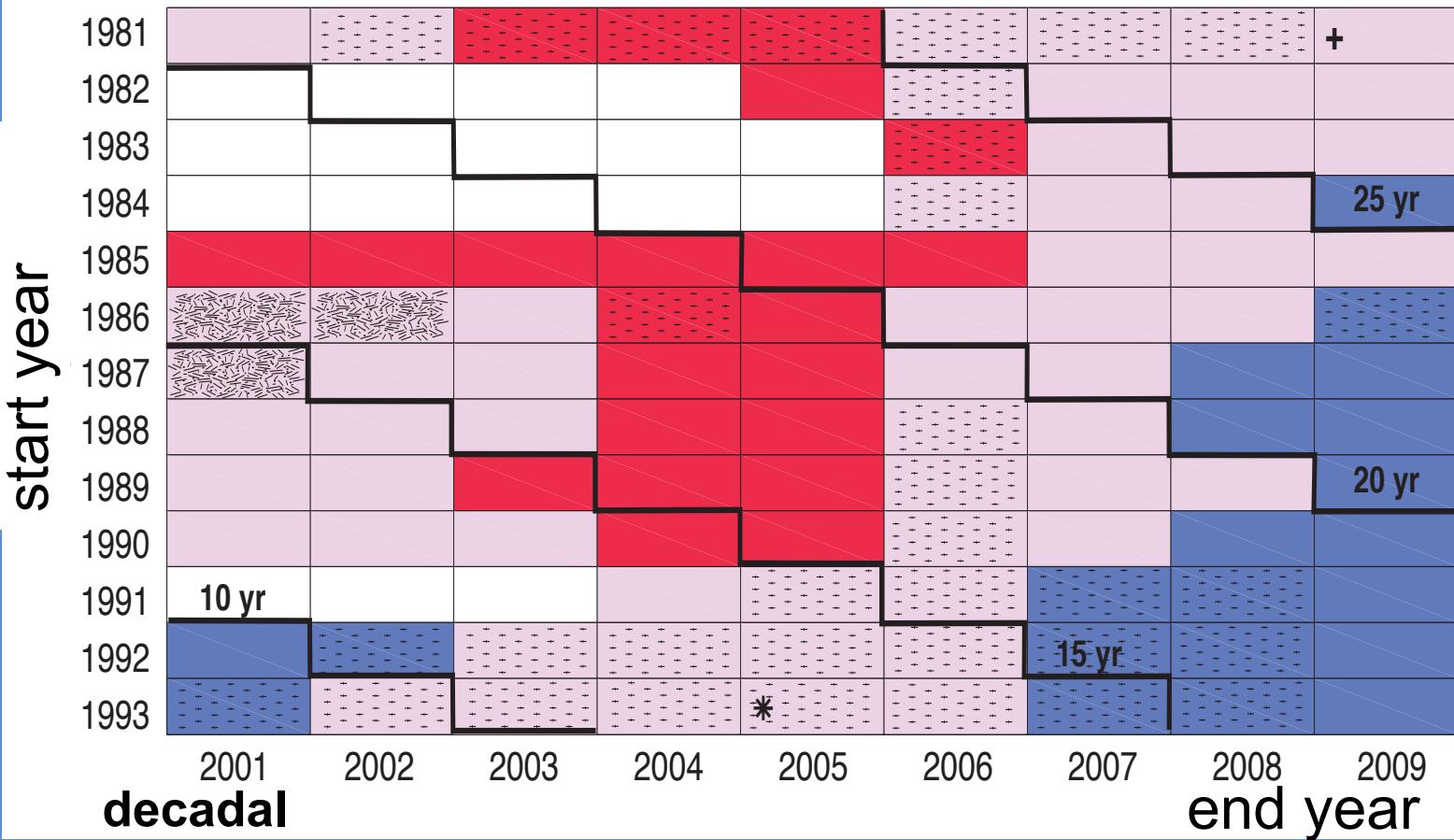
# Considering varying timescales

Dotted = warming influence significant



## SUBPOLAR BIOME (SP-SS)

multi-decadal

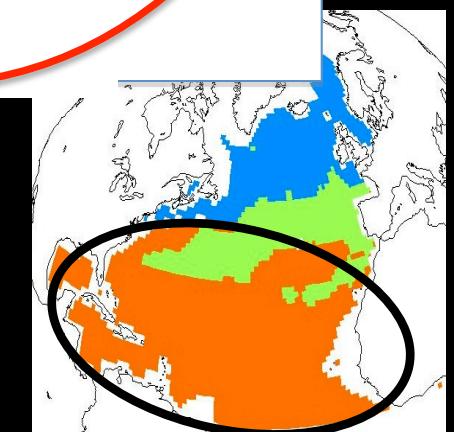
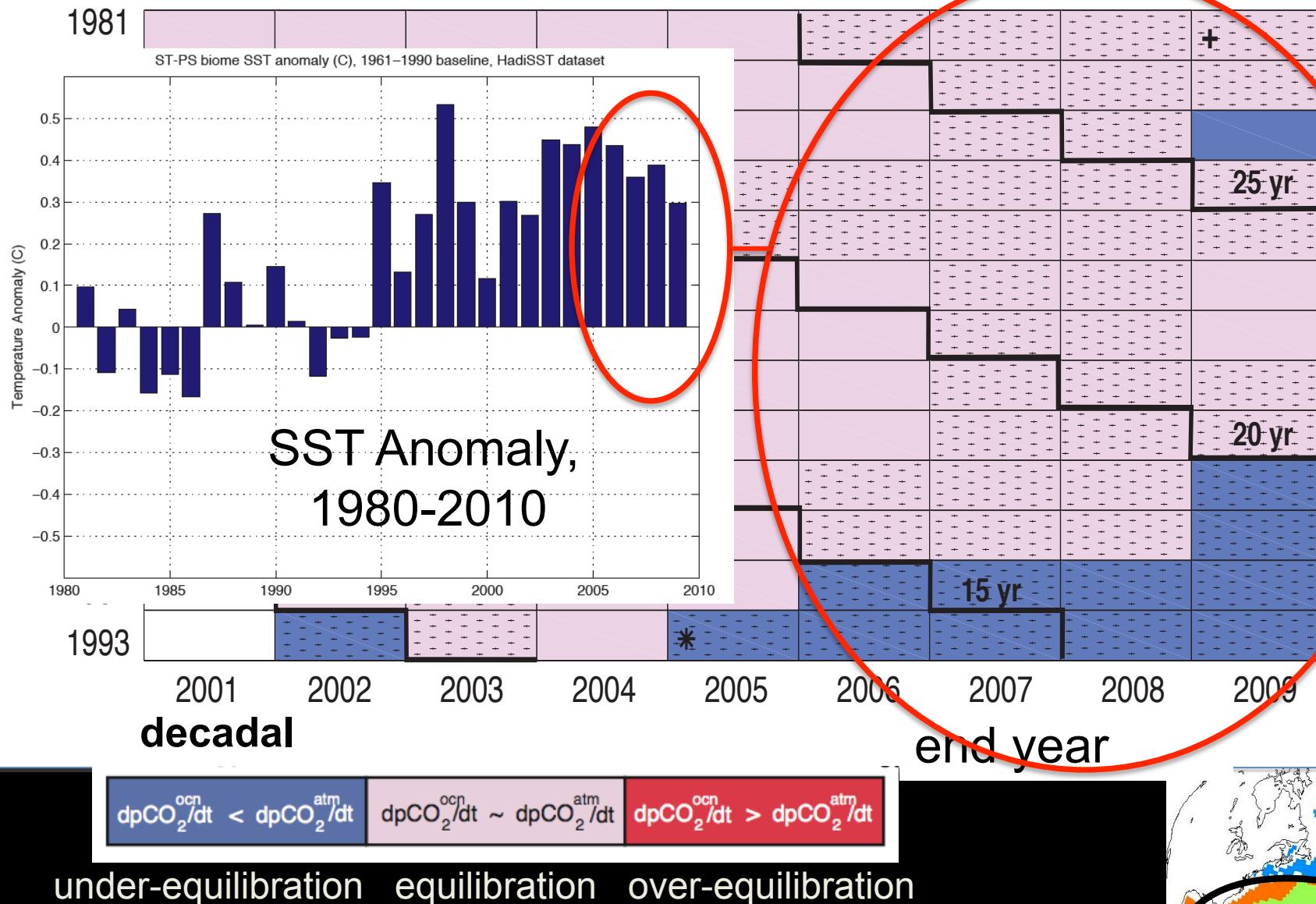


$d\text{pCO}_2^{\text{ocn}}/dt < d\text{pCO}_2^{\text{atm}}/dt$

$d\text{pCO}_2^{\text{ocn}}/dt \sim d\text{pCO}_2^{\text{atm}}/dt$

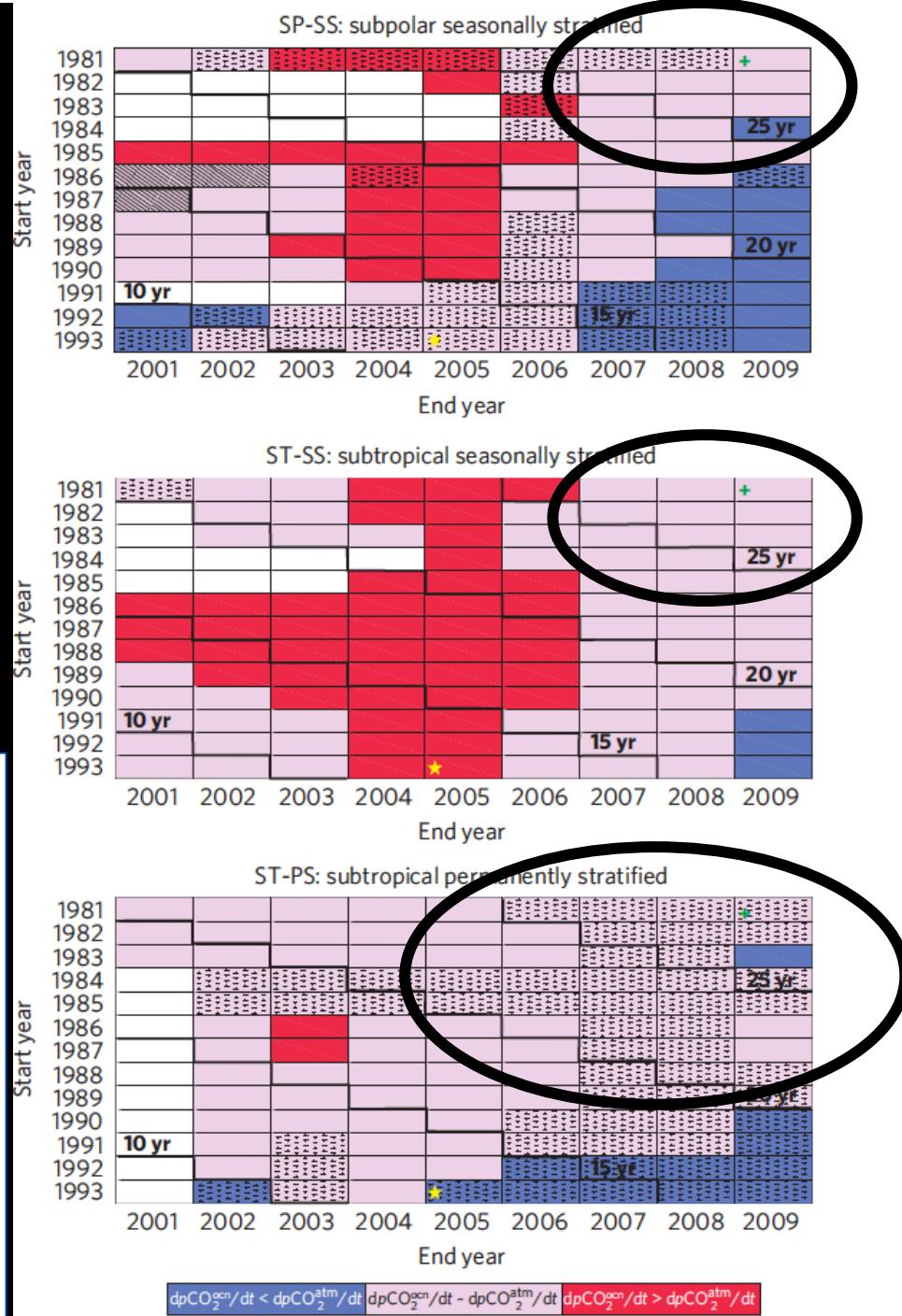
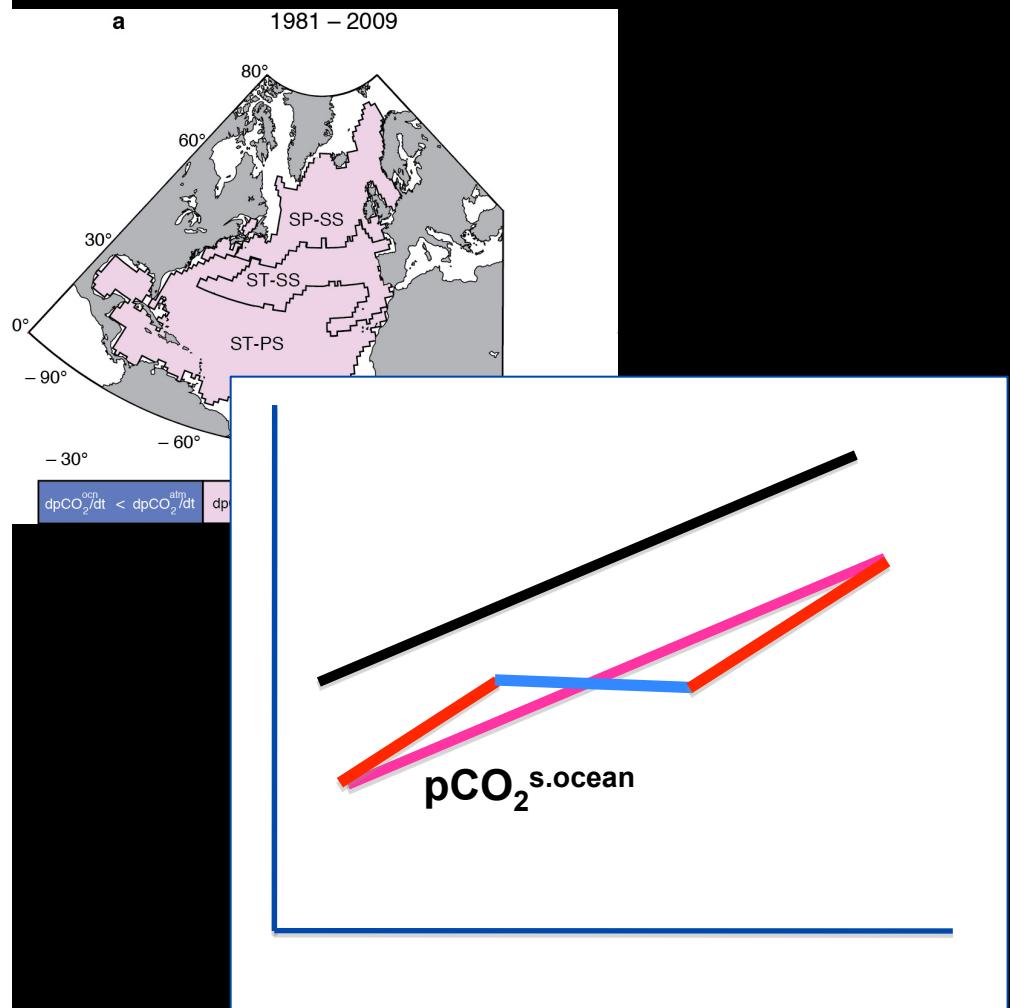
$d\text{pCO}_2^{\text{ocn}}/dt > d\text{pCO}_2^{\text{atm}}/dt$

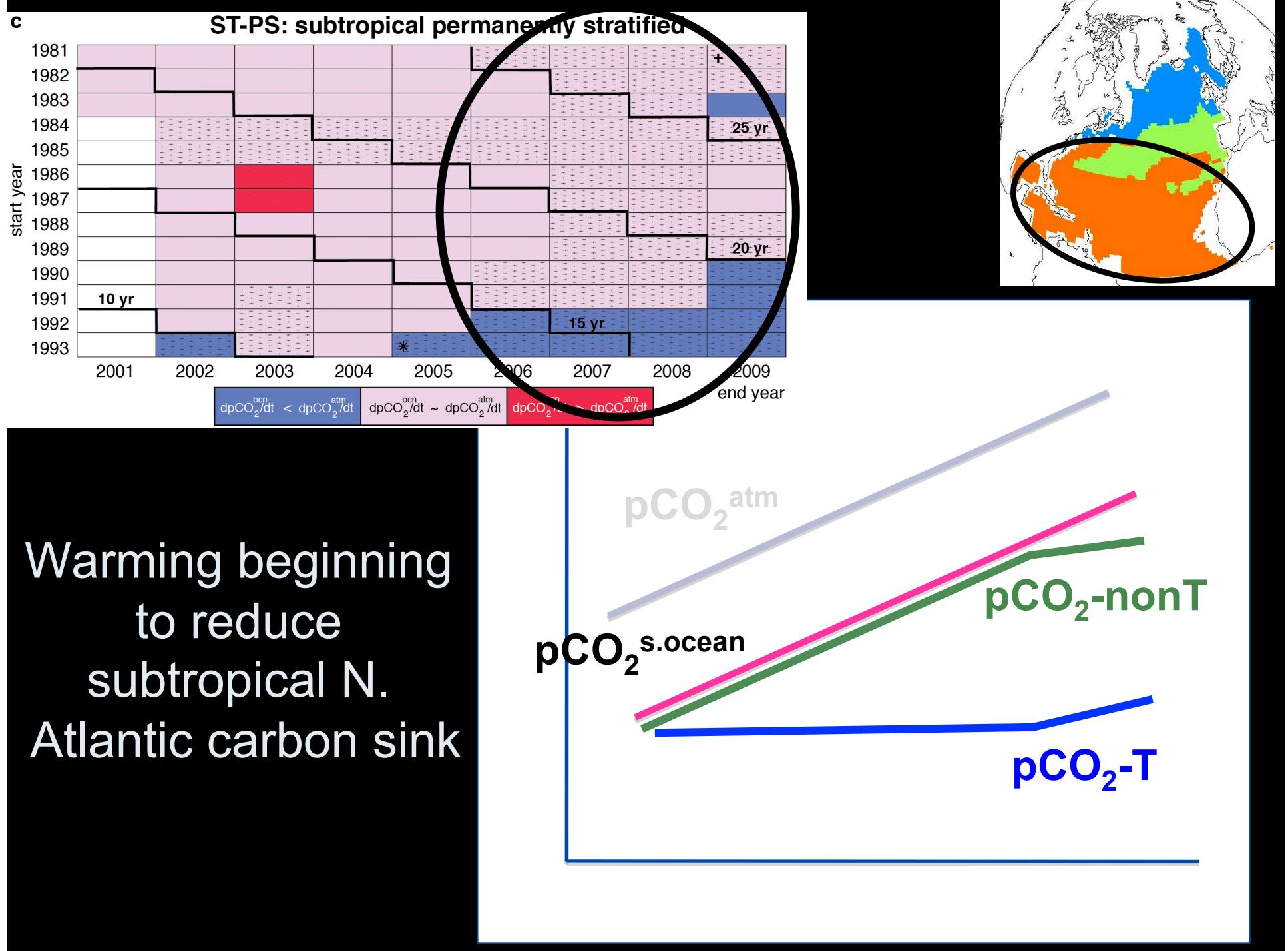
# SUBTROPICAL PERMANENTLY STRATIFIED (ST-PS)



# Conclusions: North Atlantic

# At least 20-25 years for carbon accumulation to dominate pCO<sub>2</sub><sup>s.ocean</sup> trend



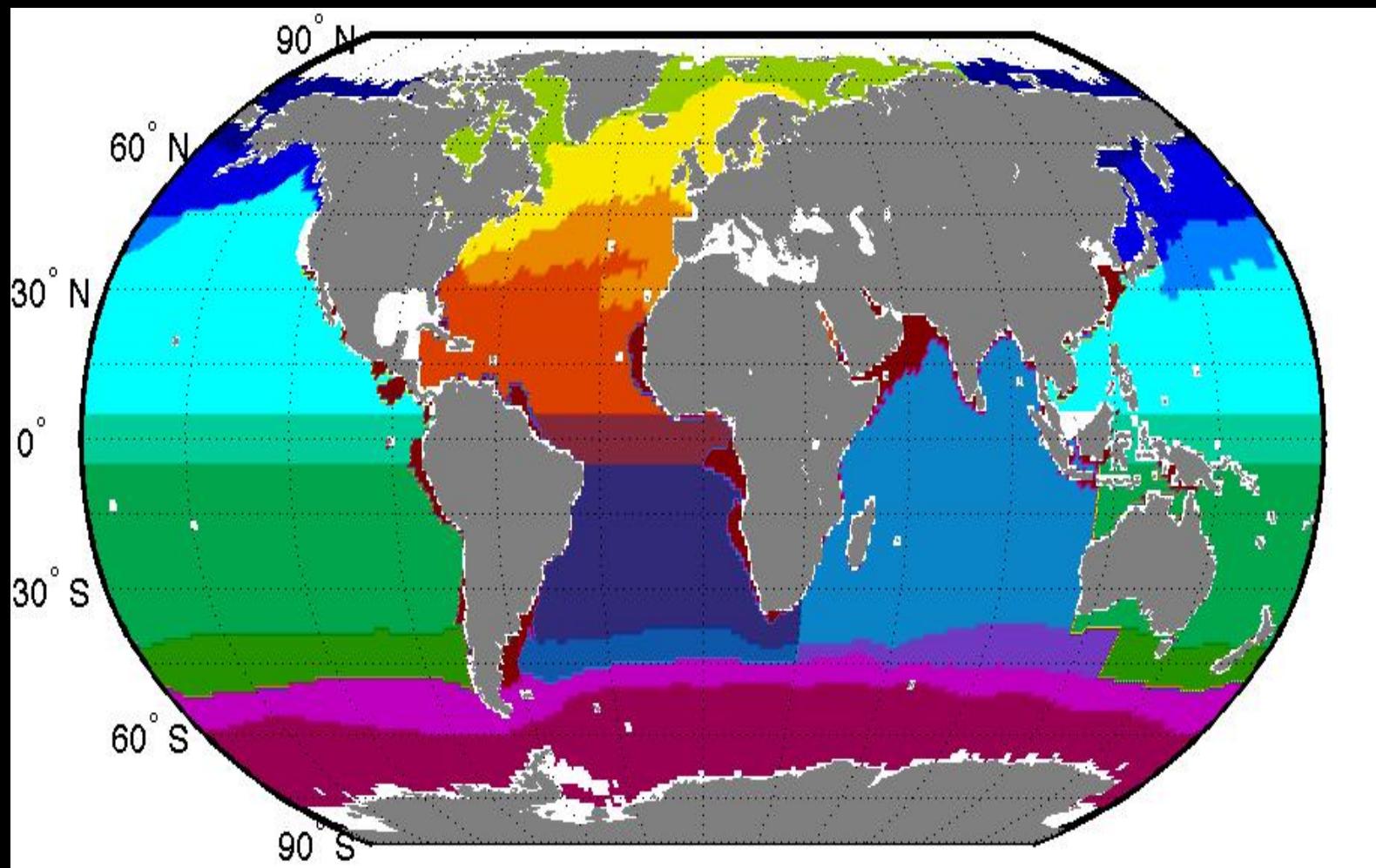


# Modeling assessment of global CO<sub>2</sub> sink impacts due to climate change: 1981-2007 = -0.20 PgC/decade

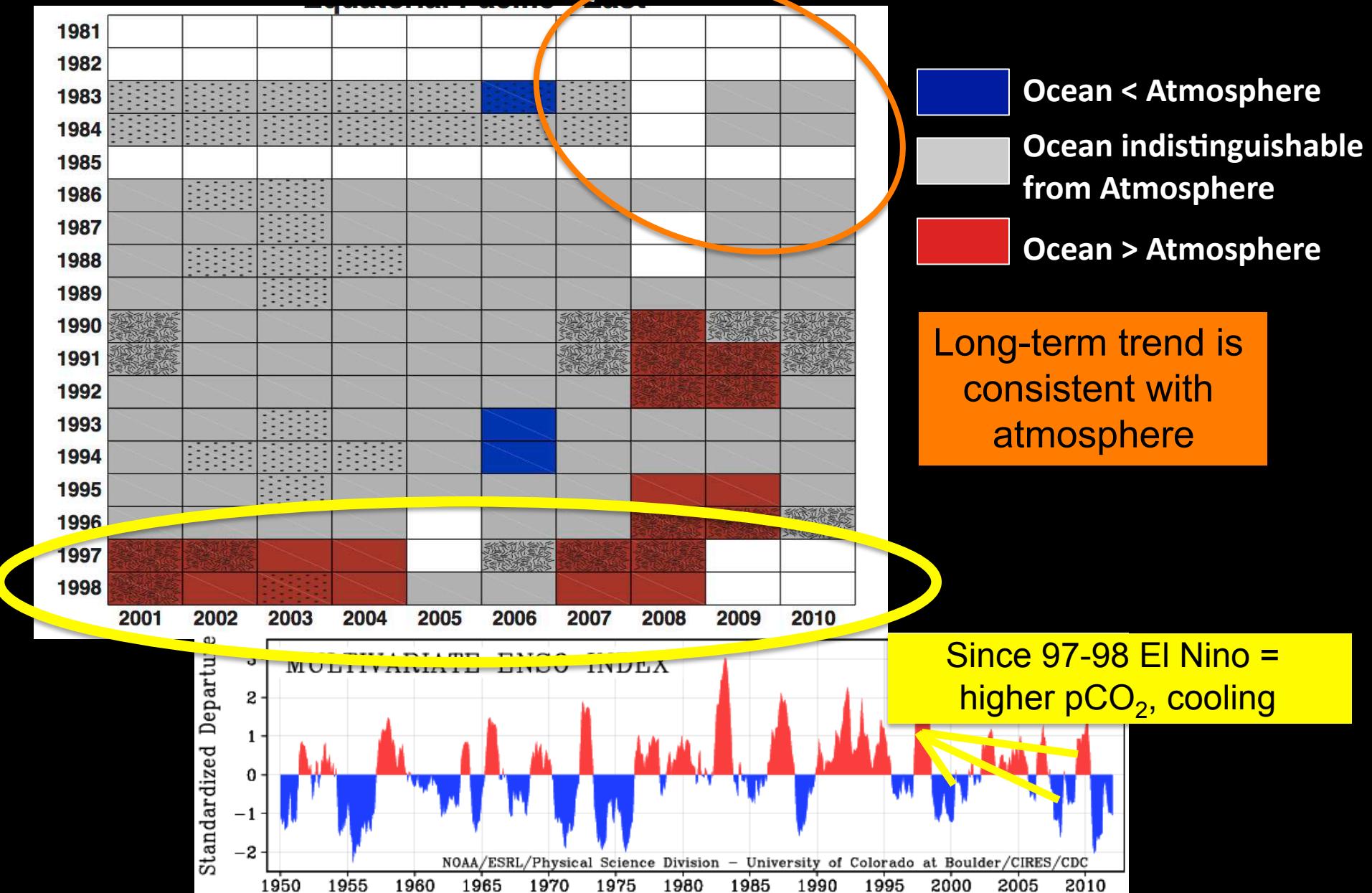
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Nonlinear	-32%	>65% in Tropics

LeQuéré et al. 2010

# Current Work: Global Extension (also see poster by Fay)



# EASTERN EQUATORIAL PACIFIC



# Questions?

